

High temperature sensitivity of monoterpene emissions from global vegetation

Efsttrios Bourtsoukidis (✉ e.bourtsoukidis@cyi.ac.cy)

The Cyprus Institute <https://orcid.org/0000-0001-5578-9414>

Andrea Pozzer (✉ andrea.pozzer@mpic.de)

Max Planck Institute for Chemistry <https://orcid.org/0000-0003-2440-6104>

Jonathan Williams (✉ jonathan.williams@mpic.de)

Max Planck Institut für Chemie <https://orcid.org/0000-0001-9421-1703>

David Makowski (✉ david.makowski@inrae.fr)

INRAE - University of Paris-Saclay <https://orcid.org/0000-0001-6385-3703>

Josep Penuelas (✉ josep.penuelas@uab.cat)

CSIC, Global Ecology Unit CREAM-CSIC-UAB, Cerdanyola del Vallès 08193, Catalonia, Spain

<https://orcid.org/0000-0002-7215-0150>

Vaseilios Matthaios (✉ v.matthaios@bham.ac.uk)

University of Birmingham

Georgia Lazoglou (✉ g.lazoglou@cyi.ac.cy)

The Cyprus Institute

Ana Yañez-Serrano (✉ ana.yanez@idaea.csic.es)

Institute of Environmental Assessment and Water Research (IDAEA)

Philippe Ciais (✉ philippe.ciais@lsce.ipsl.fr)

Laboratoire des Sciences du Climat et de l'Environnement <https://orcid.org/0000-0001-8560-4943>

Jos Lelieveld (✉ Jos.Lelieveld@mpic.de)

Max Planck Institute for Chemistry <https://orcid.org/0000-0001-6307-3846>

Mihalis Vrekoussis (✉ mvrekous@uni-bremen.de)

University of Bremen <https://orcid.org/0000-0001-8292-8352>

Nikolaos Daskalakis (✉ daskalakis@uni-bremen.de)

University of Bremen

Jean Sciare (✉ j.sciare@cyi.ac.cy)

Article

Keywords:

DOI: <https://doi.org/>

License:  This work is licensed under a Creative Commons Attribution 4.0 International License.

[Read Full License](#)

Additional Declarations: (Not answered)

High temperature sensitivity of monoterpene emissions from global vegetation

E. Bourtsoukidis^{1,2,*}, A. Pozzer², J. Williams^{2,1}, D. Makowski³, J. Peñuelas^{4,5}, V. N. Matthaïos^{6,7}, G. Lazoglou¹, A. M. Yañez-Serrano⁸, J. Lelieveld^{2,1}, P. Ciais^{9,1}, M. Vrekoussis^{10,1,11}, N. Daskalakis¹⁰, and Jean Sciare¹

¹Climate and Atmosphere Research Center (CARE-C), The Cyprus Institute; Nicosia, Cyprus.

²Max Planck Institute for Chemistry (MPIC); Mainz, Germany.

³French National Institute for Agriculture, Food, and Environment (INRAE), University Paris-Saclay, AgroParisTech (UMR MIA); Paris, France.

⁴Center for Research Ecology and Forestry Applications (CREAF); Bellaterra, Barcelona, Spain.

⁵Global Ecology Unit CREAM-CSIC-UAB; Bellaterra, Barcelona, Spain.

⁶School of Geography, Earth and Environmental Sciences, University of Birmingham; Birmingham, UK.

⁷Department of Environmental Health, Harvard T.H. Chan School of Public Health; Boston, MA, USA.

⁸Institute of Environmental Assessment and Water Research (IDAEA); Barcelona, Spain.

⁹Laboratoire des Sciences du Climat et de l'Environnement (LSCE); Gif-sur-Yvette, France.

¹⁰Institute of Environmental Physics (IUP), University of Bremen; Bremen, Germany.

¹¹Center of Marine Environmental Sciences (MARUM), University of Bremen; Bremen, Germany.

Correspondence: e.bourtsoukidis@mpic.de

Abstract

Terrestrial vegetation emits vast amounts of monoterpenes into the atmosphere, influencing ecological interactions and atmospheric chemistry. Global emissions, mostly driven by responses to abiotic stress, are simulated as a function of temperature with a fixed exponential relationship (β coefficient) across forest ecosystems and environmental conditions. We applied meta-analysis algorithms on all published monoterpene emission data and show that this relationship is more intricate and sensitive than previously thought. We find that co-occurring environmental stresses amplify the temperature sensitivity of the emissions that is primarily related to the specific plant functional type (PFT). On average, warmer ecosystems appear more sensitive, indicating the adjustment of plants in response to thermal stress. Implementing a PFT-dependent β in a biogenic emission model, coupled with a chemistry – climate model, demonstrated that atmospheric processes are exceptionally dependent on monoterpene emissions which are subject to amplified variations under rising temperatures.

Introduction

Plants emit large amounts of monoterpenes (MT; C₁₀H₁₆) in response to abiotic drivers and biotic interactions, with significant feedbacks to the ecophysiological and biogeochemical processes of terrestrial ecosystems^{1–3}. Once released, these Volatile Organic Compound (VOC) species affect atmospheric chemistry and physics due to their high reactivity and propensity to form particles, thereby influencing air quality, atmospheric oxidative capacity and radiative balance-related climate feedbacks^{4,5}.

Considered as the ‘thermometer’ of plants⁶, monoterpene emissions from global vegetation are fundamentally simulated with an empirical approach that considers temperature as a key regulator of MT emissions by global vegetation⁷:

44 $E_{MT} = E_{30} \exp(\beta(T - T_{30}))$ [1]

45 Here, E_{30} is the emission potential under standardized conditions (30 °C), and β is an
46 empirical coefficient (°C⁻¹) derived from the fit of the regression between measured emissions
47 and temperature (T). This equation remains the best approximation for MTs stored in leaf
48 pools, but an additional light-dependent function, similar to that used for isoprene, must be
49 applied for instantaneously produced MTs⁷ (Supplementary Text).

50 To date, a fixed dependence on temperature ($\beta=0.10$ °C⁻¹) is used for all types of vegetation.
51 The purpose of this meta-analysis is to conduct a detailed evaluation of this parameter by
52 collecting and analyzing all experimentally derived β coefficients published (see
53 Supplementary Information references) since the first observation of this relationship in
54 1980⁸. Moreover, we aim to assess potential refinements of β and investigate the resulting
55 impact on atmospheric chemistry and physics by performing coupled simulations between the
56 most established emission model for biogenic trace gases (MEGAN)⁷ and an atmospheric
57 chemistry – climate model (EMAC)⁹.

58 **Results**

59 **Meta-analysis of monoterpene observations**

60 Screening of all indexed research articles and data sets published during 1980-2020 yielded a
61 collection of 696 β coefficients of regression fits obtained under diverse conditions. The
62 reported values varied widely, ranging from -0.07 to 2.39 °C⁻¹. To identify the factors that
63 explain this variability, we collected an additional 35 parameters that described the
64 experimental procedure, location, time, vegetation state (i.e. the plant functional type and tree
65 characteristics) and the characteristics of MT emission (Supplementary Table 1).

66 Almost all observations were in the Northern Hemisphere, and the majority of the studies
67 were conducted under field conditions, using branch enclosures and offline chemical analyses
68 for quantifying VOCs using adsorbent tubes (Supplementary Figs. 1-2). The derivation of the
69 rate of emission (enclosures vs eddy covariance) and the techniques of sampling and
70 analyzing VOCs did not differ β amongst the studies (offline vs. online). Attempts to establish
71 clear relationships with the potential drivers were hindered by the complexity of inter-related
72 parameters and in some cases the lack of reported information for each parameter.

73 The clearest observed pattern was the increase in β with the coefficient of determination (R^2)
74 from the regression fits between emission and temperature. Our meta-analysis, based on a
75 standard random-effect statistical model¹⁰, determined the values of β within 95% confidence
76 intervals, across R^2 bins (Fig. 1). We identified a relationship between the quality of
77 regression fit and β for MT emissions that had larger values as the quality of the fit increased.
78 The majority of the data (ca. 52%) had an $R^2 > 0.6$, and the reported dependencies on
79 temperature were significantly higher than the established value of 0.10 °C⁻¹. This observation
80 indicated a clear underestimation of a globally uniform β . Interestingly, β obtained from field
81 data was systematically higher than β obtained under controlled conditions, suggesting
82 amplified sensitivities to temperature under natural conditions, when environmental stresses
83 co-occur.

84 A low R^2 indicates a weak to non-existent relationship between MT emissions and
85 temperature, so we chose to disregard the lowest 10% of the reported β (56 values), which
86 represented $R^2 < 0.2$. This approach increased the quality of our data set while retaining a
87 substantial number of observations for further statistical analyses. R^2 was not reported in 81
88 cases, so the overall data that we considered for further analysis are comprised of 559
89 experimental values.

90 We applied a machine learning analysis using the feature-importance ranking (FIR) approach,
91 with individual parameters for vegetation, location, and time to better identify the most
92 important parameters that influenced the β values (Fig. 2). The results indicated that the type
93 of ecosystem, particularly the specific plant functional type (PFT; as defined in Guenther et
94 al.⁷), was important to the derived temperature sensitivities. The geographic location (latitude
95 and longitude) remained associated with the vegetation types that grow within the
96 corresponding geographically distributed ecosystems. Our results nonetheless indicated that
97 the responses of MT emissions to temperature could vary dynamically across the life span of
98 plants. Finally, the FIR analysis did not clearly identify seasonality or a relationship between
99 tree height and the coefficient of temperature dependency.

100 **Revised temperature-dependent coefficients (β)**

101 The temperature-dependent coefficients (β) were used for each PFT in MEGAN⁷ (Fig. 3).
102 Surprisingly, we did not find any β reported for agricultural ecosystems or grasslands. Almost
103 all PFTs (other than boreal needle leaf deciduous trees and broadleaf deciduous shrubs) had
104 values significantly larger than $0.10 \text{ }^\circ\text{C}^{-1}$. This finding was particularly important for the
105 widespread boreal needle leaf evergreen forests ($\beta_{\text{NEB}} = 0.15 \text{ }^\circ\text{C}^{-1}$; range $0.12\text{-}0.19 \text{ }^\circ\text{C}^{-1}$) and
106 tropical broadleaf evergreen forests ($\beta_{\text{BETf}} = 0.20 \text{ }^\circ\text{C}^{-1}$; range $0.14\text{-}0.3 \text{ }^\circ\text{C}^{-1}$). Compared to
107 needle leaf trees and broadleaf shrubs, broadleaf trees display higher dependencies on
108 temperature in the release of MTs by their foliage. We note that removing the lowest 10%
109 from our dataset had negligible impact on the revised values (Supplementary Fig. 3). A
110 uniform β was calculated from all 559 data points as $0.13 \pm 0.01 \text{ }^\circ\text{C}^{-1}$, 30% higher than the
111 established value.

112 A β coefficient for each MT species was derived, taking into account the importance of
113 processes of atmospheric oxidation due to their different rates of reaction with atmospheric
114 radicals¹¹. No significant differences were identified across all MTs, even though β was
115 higher for Δ^3 -carene and trans- β -ocimene (Supplementary Fig. 4).

116 Summarizing the reconsiderations needed for determining β for the MTs, we concluded that
117 the most sensible approach was to assign a different value for each PFT. Alternatively, a
118 uniform value of $0.13 \text{ }^\circ\text{C}^{-1}$ could be considered, and potential discrepancies for different
119 monoterpenes may not be disregarded due to insufficient published data.

120 **Model simulations**

121 To comprehensively evaluate the implications of revised MT emission temperature
122 dependencies, we performed coupled simulations employing the most widely used VOC
123 emission model (MEGAN)⁷ and a global atmospheric chemistry – climate model (EMAC)⁹.
124 The simulations were performed at $1.8 \text{ }^\circ\text{N}$ by $1.8\text{-}^\circ\text{E}$ resolution and with hourly outputs for

125 the original version of the model ($\beta=0.10\text{ }^{\circ}\text{C}^{-1}$) and three additional scenarios for comparison.
126 First, the MEGAN code was modified (“PFT” simulation; Supplementary Code 1) to assign
127 a different β for each PFT. For the PFTs that do not have any measurements, we used the
128 value of $0.13\text{ }^{\circ}\text{C}^{-1}$ that considers the complete dataset (Fig. 3). Second, β was uniformly set to
129 $0.13\text{ }^{\circ}\text{C}^{-1}$, similar to the current approach (“ALL” simulation). Third, β was modified based
130 on β derived for each MT (“MTRP” simulation). All three simulations produced significant
131 and similar changes in the rate of emission and atmospheric feedbacks compared to the
132 current version. We focused on the most sensible approach (considering Figs. 2, 3 and
133 Supplementary Figs. 3-4), which was the revision of β across PFTs.

134 The global annual average of MT emissions was 13% lower (Supplementary Table 2). The
135 most significant changes were found for Siberia and were particularly pronounced during
136 winter. We attributed this large discrepancy to the lower β from broadleaf deciduous shrubs
137 and the generally low MT emissions in this region. Average annual β was higher for tropical
138 ecosystems, particularly for the Amazon rainforest, where MT emissions were locally up to
139 ca. 30% higher. The simulations indicated that MT emissions could even double during the
140 Amazonian dry season despite the similar annual averages (Supplementary Fig. 5).
141 Investigating the daily extremes for each season indicated that MT emissions could
142 occasionally be more than three-fold higher in both tropical and arctic ecosystems
143 (Supplementary Fig. 6).

144 The most pronounced difference was the diurnal variation in the rates of emission
145 (Supplementary Figs. 7-8). Tropical forests had particularly pronounced diel cycles that could
146 be further amplified on seasonal and daily bases. Maximum differences in the daily standard
147 deviation could increase by more than 300%, demonstrating the large variations in MT
148 emissions introduced by increased dependencies on temperature. We applied Eq. 1 using data
149 for temperature collected at the Amazonian Tall Tower Observatory (ATTO) in 2014 and
150 2015¹² to demonstrate the dynamics of PFT-dependent β (Supplementary Fig. 9). The increase
151 in β (from 0.10 to $0.20\text{ }^{\circ}\text{C}^{-1}$) led to slightly lower average emissions. The variations in the
152 standard deviation of the hourly simulated emissions, however, averaged twice as large and
153 were particularly pronounced under extreme conditions such as the El Niño year of 2015.

154 The exponential relationship between temperature and MT emission is standardized at $30\text{ }^{\circ}\text{C}$,
155 and we found that significant changes in the simulated emissions occurred at either low or
156 very high temperatures (Fig. 5, Supplementary Fig. 10). We derived β weighted by area
157 coverage for each PFT to illustrate the simplified dynamics of emission for the three dominant
158 ecosystems (Supplementary Table 3). The current models of emission have clearly
159 overestimated MT emissions under $30\text{ }^{\circ}\text{C}$. In contrast, the current models have
160 underestimated the emissions for temperatures above $30\text{ }^{\circ}\text{C}$, helping to explain the observed
161 discrepancies in diurnal variation of MT emission in warm environments^{13,14}.

162 The emission potential at standard conditions (E_{30}) was another important parameter that
163 defined the rate of MT emissions. We chose to perform the simulations adopting the values in
164 MEGAN because our analysis identified large uncertainties in the potential emissions, which
165 increased further due to the uncertainties in the conversion of units (typically from $\text{ng C g}(\text{dry}$

166 weight)⁻¹ h⁻¹ to $\mu\text{g C m}^{-2} \text{ h}^{-1}$) and the leaf area index derived from the literature and satellite
167 observations (Supplementary Figs. 11-12).

168 **Atmospheric implications**

169 The direct implications on the processes of atmospheric oxidation were evaluated by
170 investigating the differences in concentration between model simulations. The hydroxyl
171 radical (OH) is the most important daytime oxidant that reacts strongly with most organic
172 molecules, particularly MTs. Our results showed that using a PFT-dependent β increases OH
173 concentrations globally (Fig. 4). The boreal-forest belt was particularly affected, with ca. 12%
174 higher annual average OH concentrations compared to the base simulation ($\beta=0.10 \text{ }^\circ\text{C}^{-1}$).
175 Increased OH concentrations in NH summer were particularly pronounced (up to 20%)
176 (Supplementary Fig. 13). Significant daytime changes in MT emissions in EMAC indicated
177 that OH concentrations could occasionally be even as much as double compared to what is
178 currently simulated (Supplementary Fig. 14).

179 In contrast to OH, average tropospheric concentrations of ozone (O₃) were only moderately
180 affected as the maximum O₃ concentration change in the summer was within 4 - 5% for the
181 Northern Hemisphere (Supplementary Fig. 15). Regional changes in O₃ concentration on
182 specific days, however, could increase by up to 10% (Supplementary Fig. 16). Monoterpenes
183 contributed both to the production (in the presence of NO_x) and chemical removal of
184 tropospheric O₃ by reaction schemes that could differ considerably between locations (e.g.
185 VOC or NO_x limited environments¹⁵). By reducing MT emissions in our updated emission
186 model, O₃ loss was generally minimized, increasing its abundance globally, except for the hot
187 and dry season of the Amazon rainforest when the high ambient MT concentrations remove
188 O₃ at higher rates compared to the base model simulations.

189 Changes in the dynamics of atmospheric oxidation in forests influenced the yield of secondary
190 organic aerosol (SOA) particles around the globe (Fig 4, Supplementary Fig. 17). The revised
191 simulations indicated lower SOA production over boreal and temperate forests. Seasonal
192 averages nonetheless indicated increased SOA formation over both tropical forests and
193 oceans. Another consistent feature was the increase in SOAs in Siberia of the order of 10%
194 considering the annual averages.

195 Our results thus underscore that the temperature dependency of MT emissions is a sensitive
196 coefficient that introduces important feedbacks in atmospheric oxidation processes and
197 reaction products. The feedbacks include an amplified seasonality and demonstrate the
198 importance of accurately simulating the temperature effects on MT emissions.

199 **Discussion**

200 As global temperatures are rising, the emissions of monoterpenes from vegetation are
201 expected to increase exponentially, adding greater amounts of these reactive species to the
202 atmosphere. The implications range from changes in crucial ecological functions (e.g. plant
203 defense and signaling)^{16,17} to atmospheric composition and eventually radiative transfer and
204 the properties of clouds⁴. Therefore, the accurate simulation of MT emissions is essential for
205 evaluating plant responses and the respective feedbacks and improving future projections.

206 Various algorithms for simulating MT emissions have been developed accounting for either
207 short-term volatilization (emission-based)^{7,8} or long-term production of MT that are linked to
208 photosynthesis (production-based)^{18,19}. The emission-based models, such as MEGAN, have
209 been widely applied to describe observations at the scales of leaves, branches, and canopies
210 by simulating global MT emissions from terrestrial vegetation, accounting for multiple
211 parameters such as temperature, light, foliar age, soil moisture concentration, and CO₂
212 inhibition⁷. Our meta-analysis focused on the effect of temperature and in particular, on the
213 evaluation of the emission-response defining temperature sensitivity (β) of MT emissions.

214 The β coefficient for the temperature-dependent term in Eq. 1 can vary across seasons^{20,21},
215 tree species²², tree age²³, developmental stage²⁴, ambient humidity²⁵, level of
216 photosynthetically active radiation²⁶, CO₂ abundance²⁷, soil moisture concentration²⁸, and
217 other drivers²⁹. Both the FIR analysis and simplified correlation analyses on the new data
218 collected in this study demonstrated that the vegetation type was the most important regulator
219 of the sensitivity to temperature of MT emissions from vegetation. On average, plants that
220 grow in warmer ecosystems appear more sensitive, indicating the adjustment of plants in
221 response to thermal stress. Applying a PFT-dependent β helps to explain the frequently
222 overestimated MT emissions in several environments^{13,30–33}.

223 Besides β , other parameters also contribute to the emissions of MT by terrestrial vegetation.
224 The potential emission at standard condition (E_{30}) is an important parameter that regulates the
225 rate of emission by forests. Our meta-analysis collected E_{30} values that vary by orders of
226 magnitude for each ecosystem (Supplementary Figs. 11-12) revealing substantial uncertainties
227 in the magnitude of simulated emissions. Furthermore, uncertainties in the leaf area index and
228 the lack of studies over widespread ecosystems including agricultural and grasslands add to
229 the unknowns and shape the direction of future research. Our results demonstrated that co-
230 occurring environmental stresses amplify the temperature responses (β) of MT emissions. The
231 accurate formulation of a parameterization that incorporates diverse and frequently co-
232 occurring drivers of emission is challenging, because plants from different or even the same
233 ecosystems have developed different strategies to cope with environmental stresses. Plants are
234 also living organisms, so they are further subject to interspecific variability of emissions of
235 different MTs³⁴.

236 Relating these uncertainties to atmospheric processes, we identified important implications on
237 the mechanisms of atmospheric oxidation and consequently SOA formation. Introducing PFT-
238 dependent β decreased MT emissions under mean temperatures below 30 °C, thereby
239 increasing global OH concentrations. A critical review of tropospheric OH radicals identified
240 significant and unexplained discrepancies between measured and simulated OH
241 concentrations in unpolluted forested regions³⁵. By revising MT emissions through their
242 temperature sensitivity, our updated simulations help explain these observed discrepancies,
243 corroborating the strong links between the emissions and atmospheric processes.

244 With newly discovered biogenic MT sources^{12,36–38} and the challenge of modeling co-
245 occurring environmental stresses on the biosphere, our results demonstrate the need for more
246 process-oriented research of biosphere-atmosphere interactions, particularly in tropical,
247 grassland and agricultural ecosystems. As climatic warming intensifies, biogenic VOC

248 emissions from global forests will play a crucial role in the evaluation of the health of
249 ecosystems, and influence the atmospheric oxidation capacity with implications for the
250 chemical composition, aerosols and climate.

251 **Methods**

252 **Experimental design**

253 Searching the Web of Science, selecting results from the WOS, BCIBIOSIS, CCC, DRCI, and
254 RSCI databases, and using the keywords “monoterpenes”, “emissions”, and/or
255 “temperature” identified 745 peer-reviewed studies published between 1980 and 2020.
256 Screening of these articles (page by page) decreased the number of studies to 84 that reported
257 β derived from regression fits of experimental observations. The data for tropical plants were
258 not sufficient to perform reliable statistical analyses, so three additional studies (which
259 reported only data) were considered and analyzed. The timelines for the data for temperature
260 and MT emission were extracted from published plots^{6,39} using a web-based tool
261 (WebPlotDigitizer; <https://automeris.io>), and data were directly obtained from Yáñez-Serrano
262 et al.⁴⁰. We compiled a data set consisting of 696 values of β . To the best of our knowledge,
263 we accounted for all values reported in the literature. We may, however, have missed some,
264 mainly because they were not appropriately indexed in the literature data bases.

265 We investigated potential relationships with experimental, geographical, plant-specific,
266 seasonal, and regression-fit variables by extracting all available information for 35 parameters
267 (Supplementary Table 1), vectorising them (i.e. annotated/assigned a number to each
268 character class), and then proceeded with our statistical analyses.

269 The main conclusion of each study was classified in three categories: those driven by
270 temperature, temperature and light, and temperature and/or other environmental parameters
271 (e.g. soil moisture concentration and relative humidity). The experimental approaches used
272 varied in environmental conditions, determination of the rate of emission, VOC sampling, and
273 chemical analysis (Supplementary Fig. 2). These variables were comprehensively reported in
274 all studies, so the annotation was relatively straightforward. The β data points were annotated
275 using the method that was used to collect the majority of the data only when two sampling
276 methods were used. The regression fits for individual monoterpenes were annotated
277 accordingly when measurements were performed using gas chromatography.

278 The year, month, and season of the experiments were assessed. The experiments were
279 typically conducted within one year. If the experiments were conducted over two consecutive
280 years and if the regression fits were applied for all available data, we used the year in which
281 the majority of the observations were collected. If the experiments were conducted in the
282 same month of two different years, then the first year was annotated in the data points. If the
283 year of the experiments was not reported (typical for laboratory experiments), then we
284 annotated the year prior to publication. The month of an experiment was annotated with
285 numbers 1-12, and 13 was used for longer and/or mixed periods across years. Seasonal
286 observations and laboratory experiments were annotated differently. The seasons were
287 annotated accordingly in this classification.

288 The exact location of the field experiments included the latitude, longitude, and meters above
289 sea level. If the authors reported a specific station instead of the exact coordinates, we used
290 the cited literature to derive this information, but we used Google Earth and derived the
291 coordinates if deriving this information was not possible. Information for the tree species
292 (both Latin and common names) was collected from each publication. Considering the
293 geographical information, we assigned each tree species to its plant functional type (PFT) as
294 defined in the Model of Emissions of Gases and Aerosols from Nature (MEGAN)⁷. Tree age
295 (in years), tree height (in meters), and sampling height (in meters) were rarely reported
296 together, so several data gaps could not be filled. Finally, the minimum, maximum, mean, and
297 range of atmospheric temperatures were filled as reported, but if not reported, the data were
298 collected from the publication figures when possible.

299 These studies had different scientific objectives, so the collection and classification of all 35
300 parameters was a strenuous task. All parameters were cross-checked by more than one author
301 of this study to avoid mistakes during collection.

302 **Statistical analysis**

303 Data were analyzed using random-effects models, which are commonly used in meta-
304 analyses^{41–43}. A first model without a covariate was fitted to estimate the global mean β across
305 all studies. Another model with functional type as a categorical covariate was then fitted to
306 the data set for analyzing the relationship between β and PFT. Finally, a categorical variable
307 including categories of R^2 covering intervals of 0.1 was defined and included as a covariate in
308 another random-effect model for studying the relationship between β and R^2 . This model was
309 fitted to the data collected under either field or controlled conditions for analyzing the
310 sensitivity of β to the type of experiment. A random study effect was included in all fitted
311 models to account for the heterogeneity between studies. The models were fitted using the
312 lmer function of the lme4 R package by restricted maximum likelihood⁴⁴. The accuracy of
313 estimated β was assessed by calculating 95% confidence intervals. All calculations were
314 implemented using R v4.1.2.

315 **Unit conversions for potential emission at standardized conditions (30 °C, E₃₀).**

316 The rates of emission of monoterpenes were typically reported in units of carbon mass per
317 gramme of foliar dry weight per hour ($\text{ng C g(dry weight)}^{-1} \text{h}^{-1}$). The biogenic emission factors
318 in MEGAN are in $\mu\text{g m}^{-2} \text{h}^{-1}$, so we used the leaf area index (LAI) and the specific leaf area
319 (SLA) for each point to convert the reported values. SLA was obtained from the TRY
320 database⁴⁵ (<https://try-db.org>, last accessed 8 June 2022). The data sets used were 1-km⁴⁶
321 maps scaled up from trait data measured in situ. The LAI data we used were the monthly
322 climatological data from the ORNL DAAC database⁴⁷. The data for SLA and LAI were
323 extracted from the above data sets, for the longitudes and latitudes of our data points while in
324 case of missing values, the nearest neighbor with valid data was used.

325 **Machine learning**

326 We used machine learning methods to examine the relationship and the importance of β
327 coefficient to the three broad categories (vegetation type, location and time variables)
328 identified in our meta-analysis. As each of these broad categories had sub-divisions, we

329 examined them using the Featured Importance Ranking (FIR) approach of machine learning.
330 Feature (or variable) importance ranking refers to a task that measures the contributions of
331 individual input features (variables) to the performance of a supervised learning model⁴⁸.
332 Feature importance ranking has become one of the most powerful tools in
333 explainable/interpretable models to facilitate understanding and discovery of key factors in a
334 specific domain^{49,50}.

335 Specifically, we used Caret and Cubist packages in R with tuning two hyper-parameters:
336 neighbors (#Instances) and committees (#Committees), which are the ones to most likely have
337 the largest effect on the final performance of the Cubist model⁵¹. Cubist is a rule-based tree
338 algorithm, where a tree is grown, and the terminal leaves contain regression models. These
339 models are based on the predictors used in previous splits. In these algorithms, the prediction
340 is made using the linear regression model at the terminal node of the tree but is smoothed by
341 taking into account the prediction from another model in the previous node of the tree (which
342 also occurs recursively up the tree). The tree is reduced to a set of rules, which initially are
343 paths from the top of the tree to the bottom⁵². The performance is taken from each
344 combination of the hyper-parameters tuning with the grid search method with cross-validation
345 (CV)⁵³. To avoid bias in data selection, we applied 10-fold CV^{53,54}, while the model's final
346 performance against the test dataset was validated using R^2 and Root Mean Square Error
347 (RMSE). We also tested an ANN and a random forest algorithm but Cubist had the smaller
348 RMSE and the greater R^2 out of the three (Supplementary Fig. 18).

349 **Model simulations**

350 We used the global ECHAM/MESSy Atmospheric chemistry – Climate (EMAC) model,
351 which simulates atmospheric chemical and meteorological processes and interactions with
352 oceans and the biosphere^{55,56}. The model uses the second version of the Modular Earth
353 Submodel System (MESSy2) to link multi-institutional computer codes. The core
354 atmospheric model was the 5th generation European Centre/Hamburg general circulation
355 model (ECHAM5), into which updates and improvements in boundary layer, radiation, and
356 convection routines have been introduced⁵⁶⁻⁵⁸. Additional descriptions, references, and
357 information for the model are available at <https://www.messy-interface.org>. We applied
358 EMAC (ECHAM5 version 5.3.02, MESSy version 2.55.0) with a spherical truncation of T63
359 (corresponding to a quadratic Gaussian grid of approximately 1.8×1.8 degrees of latitude and
360 longitude) with 31 levels of vertical hybrid pressure to 10 hPa.

361 The various submodels represented tropospheric processes and their interactions with oceans,
362 land, and human influences describing emissions, including isotopic composition, radiative
363 processes, atmospheric multiphase chemistry, aerosols, and mechanisms of deposition^{57,59}.
364 The set-up used in this simulation was identical to that used by Pozzer et al.⁹, where a detailed
365 scheme of the oxidation of volatile organic compounds (VOCs) (the Mainz Organic
366 Mechanism) was coupled to a base set of volatility (ORACLE⁶⁰) to simulate the partitioning
367 of organic gases and aerosols in unprecedented detail for a chemistry – climate model. We
368 only briefly summarized the most important characteristics (see Pozzer et al.⁹ for more
369 details).

370 EMAC simulates gas-phase and heterogeneous chemistry using the MECCA submodel, which
371 accounts for the photochemical oxidation of natural and anthropogenic VOCs⁶⁰⁻⁶². Processes
372 of aerosol microphysics and gas/aerosol partitioning were simulated using the GMXe
373 submodel⁶³. The distribution of aerosol sizes was described using seven interacting lognormal
374 modes (four hydrophilic and three hydrophobic modes). The composition of aerosols was
375 uniform within each mode (internally mixed) but could vary between modes (externally
376 mixed). The four modes of hydrophilic size (nucleation, Aitken, accumulation, and coarse)
377 encompassed the spectrum of aerosol sizes. The composition of inorganic aerosols was
378 determined using the ISORROPIA-II thermodynamic equilibrium submodel⁶⁴, which
379 calculates the gas/liquid/solid equilibrium partitioning of inorganic compounds and water.
380 The components of aeolian dust can exist in the form of mineral salts in the solid phase and
381 ions in the aqueous phase⁶⁵. The composition and atmospheric evolution of organic aerosol
382 compounds were simulated using the ORACLE sub-model, which represents classes of the
383 volatility of organics by their effective saturation concentrations⁶⁶. The biogenic emissions of
384 non-methane VOCs were calculated online using MEGAN⁷. The results of the model for the
385 last decade have been extensively tested against measured data for gases and particles from
386 ground-based networks monitoring air quality and global observations from satellites⁶⁷⁻⁷⁰.

387 The dynamics were not nudged, unlike the study by Pozzer et al.⁹, but the model ran freely,
388 forced only by climatological sea-surface temperature (SST) and sea-ice coverage (SIC)
389 obtained from the ERA5 data for 2010-2019, which allowed the model to recreate a
390 climatological meteorology without strong extremes in the forcing (SST and SIC). The
391 meteorology/radiation and the chemistry were also fully decoupled, so all simulations
392 performed developed identical binary meteorology, allowing comparisons between the
393 simulations. In summary, any difference between the concentrations of MTs in the
394 simulations were only due to changes in the emissions and not in different modes of transport.

395 We performed four simulations: 1) a BASE simulation, where MEGAN was used in its
396 standard configuration ($\beta=0.10\text{ }^{\circ}\text{C}^{-1}$), 2) an ALL simulation, where MEGAN was used in its
397 standard configuration but with β set at $0.13\text{ }^{\circ}\text{C}^{-1}$, 3) an MTRP simulation, where β was based
398 on the monoterpene emitted (Supplementary Fig. 3) a PFT simulation, where β in MEGAN
399 varied based on PFT (Fig. 3, Supplementary Code 1).

400 **Data availability**

401 The meta-analysis data are available upon request to all scientists agreeing to the Cyprus
402 Institute protocol (<https://doi.org/10.XXXXX/zenodo.XXXXX; to be provided>). The software
403 code is available at the supplementary information.

404 **Competing interests**

405 The authors declare that they have no conflict of interest.

406 **Author contributions**

407 E.B. designed the research and drafted the paper. E.B., N.D., J.P., M.V., and A.M.Y-S
408 collected the data. E.B., G.L. D.M., V.N.M., and A.P. analyzed the data. E.B. P.C. J.L., J.P.,
409 A.P., J.S and J.W. performed research. All authors contributed to editing the article.

410 Acknowledgements

411 We acknowledge the EMME-CARE project from the European Union's Horizon 2020
412 Research and Innovation Programme (grant agreement No. 856612), as well as matching co-
413 funding by the Government of the Republic of Cyprus. We acknowledge the use of the HPC
414 cluster Aether at the University of Bremen, financed by DFG within the scope of the
415 Excellence Initiative. We acknowledge the support of BmBf project ATTO (01LK1602B).

416 References

- 417 1. Vickers, C. E., Gershenzon, J., Lerdau, M. T. & Loreto, F. A unified mechanism of action for
418 volatile isoprenoids in plant abiotic stress. *Nat Chem Biol* **5**, 283–291 (2009).
- 419 2. Gershenzon, J. & Dudareva, N. The function of terpene natural products in the natural world. *Nat*
420 *Chem Biol* **3**, 408–414 (2007).
- 421 3. Loreto, F., Dicke, M., Schnitzler, J.-P. & Turlings, T. C. J. Plant volatiles and the environment: Plant
422 volatiles and the environment. *Plant Cell Environ* **37**, 1905–1908 (2014).
- 423 4. Riipinen, I. *et al.* The contribution of organics to atmospheric nanoparticle growth. *Nature Geosci*
424 **5**, 453–458 (2012).
- 425 5. Yli-Juuti, T. *et al.* Significance of the organic aerosol driven climate feedback in the boreal area.
426 *Nat Commun* **12**, 5637 (2021).
- 427 6. Jardine, K. J. *et al.* Monoterpene 'thermometer' of tropical forest-atmosphere response to
428 climate warming: Monoterpene 'thermometer'. *Plant, Cell & Environment* **40**, 441–452 (2017).
- 429 7. Guenther, A. B. *et al.* The Model of Emissions of Gases and Aerosols from Nature version 2.1
430 (MEGAN2.1): an extended and updated framework for modeling biogenic emissions. *Geosci.*
431 *Model Dev.* **5**, 1471–1492 (2012).
- 432 8. Tingey, D. T., Manning, M., Grothaus, L. C. & Burns, W. F. Influence of Light and Temperature on
433 Monoterpene Emission Rates from Slash Pine. *Plant Physiol.* **65**, 797–801 (1980).
- 434 9. Pozzer, A. *et al.* Simulation of organics in the atmosphere: evaluation of EMACv2.54 with the
435 Mainz Organic Mechanism (MOM) coupled to the ORACLE (v1.0) submodel. *Geosci. Model Dev.*
436 **15**, 2673–2710 (2022).
- 437 10. Makowski, D. *et al.* Systematic review of meta-analyses to assess the impacts of farming
438 practices - A methodological framework. <https://osf.io/byuw9> (2021)
439 doi:10.31219/osf.io/byuw9.
- 440 11. Stewart, D. J. *et al.* The kinetics of the gas-phase reactions of selected monoterpenes and cyclo-
441 alkenes with ozone and the NO₃ radical. *Atmospheric Environment* **70**, 227–235 (2013).
- 442 12. Bourtsoukidis, E. *et al.* Strong sesquiterpene emissions from Amazonian soils. *Nat Commun* **9**,
443 2226 (2018).
- 444 13. Fares, S. *et al.* Observations of Diurnal to Weekly Variations of Monoterpene-Dominated Fluxes
445 of Volatile Organic Compounds from Mediterranean Forests: Implications for Regional Modeling.
446 *Environ. Sci. Technol.* **47**, 11073–11082 (2013).
- 447 14. Sarkar, C. *et al.* PTR-TOF-MS eddy covariance measurements of isoprene and monoterpene
448 fluxes from an eastern Amazonian rainforest. *Atmos. Chem. Phys.* **20**, 7179–7191 (2020).
- 449 15. Atkinson, R. Atmospheric chemistry of VOCs and NO_x. *Atmospheric Environment* **34**, 2063–2101
450 (2000).
- 451 16. Peñuelas, J. & Staudt, M. BVOCs and global change. *Trends in Plant Science* **15**, 133–144 (2010).
- 452 17. Yáñez-Serrano, A. M. *et al.* Amazonian biogenic volatile organic compounds under global change.
453 *Glob Change Biol* **26**, 4722–4751 (2020).
- 454 18. Niinemets, U., Seufert, G., Steinbrecher, R. & Tenhunen, J. D. A model coupling foliar
455 monoterpene emissions to leaf photosynthetic characteristics in Mediterranean evergreen
456 *Quercus* species. *New Phytol* **153**, 257–275 (2002).
- 457 19. Grote, R. *et al.* Process-based modelling of isoprenoid emissions from evergreen leaves of
458 *Quercus ilex* (L.). *Atmospheric Environment* **40**, 152–165 (2006).

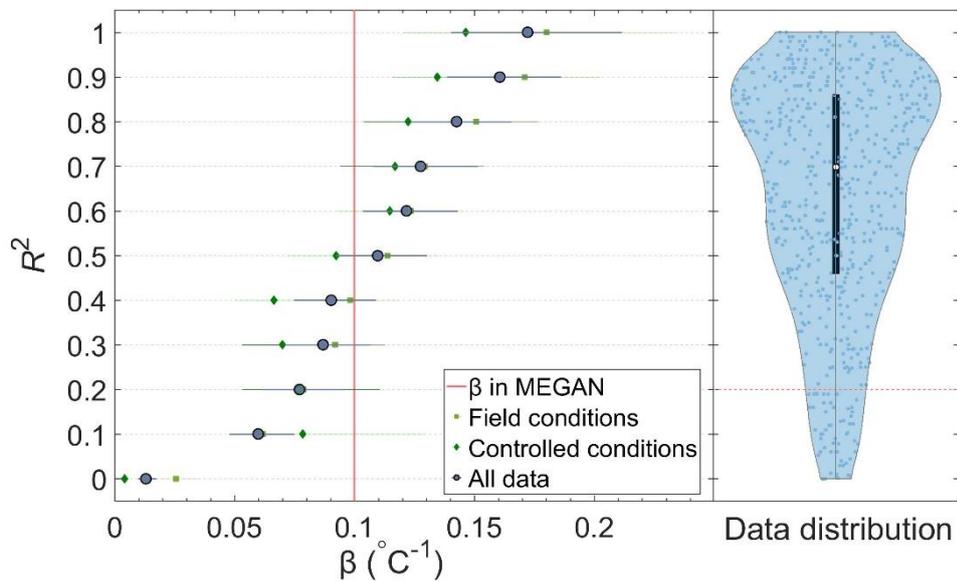
- 459 20. Bourtsoukidis, E., Williams, J., Kesselmeier, J., Jacobi, S. & Bonn, B. From emissions to ambient
460 mixing ratios: online seasonal field measurements of volatile organic compounds over a Norway
461 spruce-dominated forest in central Germany. *Atmos. Chem. Phys.* **14**, 6495–6510 (2014).
- 462 21. Tarvainen, V. *et al.* Temperature and light dependence of the VOC emissions of Scots pine.
463 *Atmos. Chem. Phys.* **5**, 989–998 (2005).
- 464 22. Fares, S. *et al.* Biogenic emissions from Citrus species in California. *Atmospheric Environment* **45**,
465 4557–4568 (2011).
- 466 23. Lim, J.-H. *et al.* Seasonal variations of monoterpene emissions from *Pinus densiflora* in East Asia.
467 *Chemosphere* **73**, 470–478 (2008).
- 468 24. Hakola, H., Rinne, J. & Laurila, T. The hydrocarbon emission rates of tea-leafed willow (*Salix*
469 *phylicifolia*), silver birch (*Betula pendula*) and European aspen (*Populus tremula*). *Atmospheric*
470 *Environment* **32**, 1825–1833 (1998).
- 471 25. Schade, G. W., Goldstein, A. H. & Lamanna, M. S. Are monoterpene emissions influenced by
472 humidity? *Geophys. Res. Lett.* **26**, 2187–2190 (1999).
- 473 26. Staudt, M. *et al.* Seasonal and diurnal patterns of monoterpene emissions from *Pinus pinea* (L.)
474 under field conditions. *Atmospheric Environment* **31**, 145–156 (1997).
- 475 27. Laffineur, Q. *et al.* Impact of diffuse light on isoprene and monoterpene emissions from a mixed
476 temperate forest. *Atmospheric Environment* **74**, 385–392 (2013).
- 477 28. Seco, R. *et al.* Ecosystem-scale volatile organic compound fluxes during an extreme drought in a
478 broadleaf temperate forest of the Missouri Ozarks (central USA). *Glob Change Biol* **21**, 3657–
479 3674 (2015).
- 480 29. Filella, I., Wilkinson, M. J., Llusà, J., Hewitt, C. N. & Peñuelas, J. Volatile organic compounds
481 emissions in Norway spruce (*Picea abies*) in response to temperature changes. *Physiol Plant* **130**,
482 58–66 (2007).
- 483 30. McKinney, K. A., Lee, B. H., Vasta, A., Pho, T. V. & Munger, J. W. Emissions of isoprenoids and
484 oxygenated biogenic volatile organic compounds from a New England mixed forest. *Atmos.*
485 *Chem. Phys.* **11**, 4807–4831 (2011).
- 486 31. Geron, C. D., Daly, R. W., Arnts, R. R., Guenther, A. B. & Mowry, Fred. L. Canopy level emissions
487 of 2-methyl-3-buten-2-ol, monoterpenes, and sesquiterpenes from an experimental *Pinus taeda*
488 plantation. *Science of The Total Environment* **565**, 730–741 (2016).
- 489 32. Tang, J. *et al.* Challenges in modelling isoprene and monoterpene emission dynamics of Arctic
490 plants: a case study from a subarctic tundra heath. *Biogeosciences* **13**, 6651–6667 (2016).
- 491 33. Warneke, C. *et al.* Biogenic emission measurement and inventories determination of biogenic
492 emissions in the eastern United States and Texas and comparison with biogenic emission
493 inventories. *J. Geophys. Res.* **115**, D00F18 (2010).
- 494 34. Bäck, J. *et al.* Chemodiversity of a Scots pine stand and implications for terpene air
495 concentrations. *Biogeosciences* **9**, 689–702 (2012).
- 496 35. Stone, D., Whalley, L. K. & Heard, D. E. Tropospheric OH and HO₂ radicals: field measurements
497 and model comparisons. *Chem. Soc. Rev.* **41**, 6348 (2012).
- 498 36. Vanhatalo, A. *et al.* Scots Pine Stems as Dynamic Sources of Monoterpene and Methanol
499 Emissions. *Front. For. Glob. Change* **2**, 95 (2020).
- 500 37. Edtbauer, A. *et al.* Cryptogamic organisms are a substantial source and sink for volatile organic
501 compounds in the Amazon region. *Commun Earth Environ* **2**, 258 (2021).
- 502 38. Mäki, M., Aalto, J., Hellén, H., Pihlatie, M. & Bäck, J. Interannual and Seasonal Dynamics of
503 Volatile Organic Compound Fluxes From the Boreal Forest Floor. *Front. Plant Sci.* **10**, 191 (2019).
- 504 39. Langford, B. *et al.* Fluxes and concentrations of volatile organic compounds from a South-East
505 Asian tropical rainforest. *Atmos. Chem. Phys.* **10**, 8391–8412 (2010).
- 506 40. Yáñez-Serrano, A. M. *et al.* Heat stress increases the use of cytosolic pyruvate for isoprene
507 biosynthesis. *Journal of Experimental Botany* **70**, 5827–5838 (2019).
- 508 41. Borenstein, M., Hedges, L. V., Higgins, J. P. T. & Rothstein, H. R. *Introduction to Meta-Analysis*.
509 (John Wiley & Sons, Ltd, 2009). doi:10.1002/9780470743386.

- 510 42. Hedges, L. V., Gurevitch, J. & Curtis, P. S. THE META-ANALYSIS OF RESPONSE RATIOS IN
511 EXPERIMENTAL ECOLOGY. *Ecology* **80**, 1150–1156 (1999).
- 512 43. Makowski, D., Piraux, F. & Brun, F. *From Experimental Network to Meta-analysis: Methods and*
513 *Applications with R for Agronomic and Environmental Sciences*. (Springer Netherlands, 2019).
514 doi:10.1007/978-94-024-1696-1.
- 515 44. Bates, D., Mächler, M., Bolker, B. & Walker, S. Fitting Linear Mixed-Effects Models Using **lme4**. *J.*
516 *Stat. Soft.* **67**, (2015).
- 517 45. Moreno-Martínez, Á. *et al.* A methodology to derive global maps of leaf traits using remote
518 sensing and climate data. *Remote Sensing of Environment* **218**, 69–88 (2018).
- 519 46. Camps-Valls, G. *et al.* Global trait maps at 1km resolution. (2020) doi:10.17871/TRY.60.
- 520 47. MAO, J. & YAN, B. Global Monthly Mean Leaf Area Index Climatology, 1981-2015. 6.581285 MB
521 (2019) doi:10.3334/ORNLDAAC/1653.
- 522 48. Samek, W., Wiegand, T. & Müller, K.-R. Explainable Artificial Intelligence: Understanding,
523 Visualizing and Interpreting Deep Learning Models. *arXiv:1708.08296 [cs, stat]* (2017).
- 524 49. Matthaios, V. N. *et al.* Quantifying factors affecting contributions of roadway exhaust and non-
525 exhaust emissions to ambient PM_{10-2.5} and PM_{2.5-0.2} particles. *Science of The Total*
526 *Environment* **835**, 155368 (2022).
- 527 50. Matthaios, V. N. *et al.* Factors Influencing Classroom Exposures to Fine Particles, Black Carbon,
528 and Nitrogen Dioxide in Inner-City Schools and Their Implications for Indoor Air Quality. *Environ*
529 *Health Perspect* **130**, 047005 (2022).
- 530 51. Zhou, J. *et al.* Random Forests and Cubist Algorithms for Predicting Shear Strengths of Rockfill
531 Materials. *Applied Sciences* **9**, 1621 (2019).
- 532 52. Wang, Y. & Witten, I. *Induction of model trees for predicting continuous classes*. (University of
533 Economics, Faculty of Informatics and Statistics, 1997).
- 534 53. Kuhn, M. & Johnson, K. *Applied Predictive Modeling*. (Springer, 2013).
- 535 54. Noi, P., Degener, J. & Kappas, M. Comparison of Multiple Linear Regression, Cubist Regression,
536 and Random Forest Algorithms to Estimate Daily Air Surface Temperature from Dynamic
537 Combinations of MODIS LST Data. *Remote Sensing* **9**, 398 (2017).
- 538 55. Jöckel, P. *et al.* The atmospheric chemistry general circulation model ECHAM5/MESSy1:
539 consistent simulation of ozone from the surface to the mesosphere. *Atmos. Chem. Phys.* **6**,
540 5067–5104 (2006).
- 541 56. Jöckel, P. *et al.* Development cycle 2 of the Modular Earth Submodel System (MESSy2). *Geosci.*
542 *Model Dev.* **3**, 717–752 (2010).
- 543 57. Tost, H. *et al.* Global cloud and precipitation chemistry and wet deposition: tropospheric model
544 simulations with ECHAM5/MESSy1. *Atmos. Chem. Phys.* **7**, 2733–2757 (2007).
- 545 58. Ouwersloot, H. G., Pozzer, A., Steil, B., Tost, H. & Lelieveld, J. Revision of the convective transport
546 module CVTRANS 2.4 in the EMAC atmospheric chemistry–climate model. *Geosci. Model Dev.* **8**,
547 2435–2445 (2015).
- 548 59. Pozzer, A. *et al.* Observed and simulated global distribution and budget of atmospheric C₂-C₅
549 alkanes. *Atmos. Chem. Phys.* **10**, 4403–4422 (2010).
- 550 60. Tsimpidi, A. P., Karydis, V. A., Pozzer, A., Pandis, S. N. & Lelieveld, J. ORACLE (v1.0): module to
551 simulate the organic aerosol composition and evolution in the atmosphere. *Geosci. Model Dev.*
552 **7**, 3153–3172 (2014).
- 553 61. Sander, R. *et al.* The atmospheric chemistry box model CAABA/MECCA-3.0. *Geosci. Model Dev.* **4**,
554 373–380 (2011).
- 555 62. Lelieveld, J., Gromov, S., Pozzer, A. & Taraborrelli, D. Global tropospheric hydroxyl distribution,
556 budget and reactivity. *Atmos. Chem. Phys.* **16**, 12477–12493 (2016).
- 557 63. Pringle, K. J. *et al.* Description and evaluation of GMXe: a new aerosol submodel for global
558 simulations (v1). *Geosci. Model Dev.* **3**, 391–412 (2010).
- 559 64. Fountoukis, C. & Nenes, A. ISORROPIA II: a computationally efficient thermodynamic equilibrium
560 model for K⁺-Ca²⁺-Mg²⁺-NH₄⁺-Na⁺-SO₄²⁻-NO₃⁻-Cl⁻-H₂O aerosols. *Atmos. Chem. Phys.* **7**,
561 4639–4659 (2007).

- 562 65. Karydis, V. A., Tsimpidi, A. P., Pozzer, A., Astitha, M. & Lelieveld, J. Effects of mineral dust on
563 global atmospheric nitrate concentrations. *Atmos. Chem. Phys.* **16**, 1491–1509 (2016).
- 564 66. Tsimpidi, A. P., Karydis, V. A., Pandis, S. N. & Lelieveld, J. Global combustion sources of organic
565 aerosols: model comparison with 84 AMS factor-analysis data sets. *Atmos. Chem. Phys.* **16**,
566 8939–8962 (2016).
- 567 67. Pozzer, A. *et al.* Distributions and regional budgets of aerosols and their precursors simulated
568 with the EMAC chemistry-climate model. *Atmos. Chem. Phys.* **12**, 961–987 (2012).
- 569 68. Pozzer, A. *et al.* Effects of business-as-usual anthropogenic emissions on air quality. *Atmos.*
570 *Chem. Phys.* **12**, 6915–6937 (2012).
- 571 69. Lelieveld, J., Evans, J. S., Fnais, M., Giannadaki, D. & Pozzer, A. The contribution of outdoor air
572 pollution sources to premature mortality on a global scale. *Nature* **525**, 367–371 (2015).
- 573 70. Karydis, V. A. *et al.* Global impact of mineral dust on cloud droplet number concentration. *Atmos.*
574 *Chem. Phys.* **17**, 5601–5621 (2017).
- 575

576 **Figures**

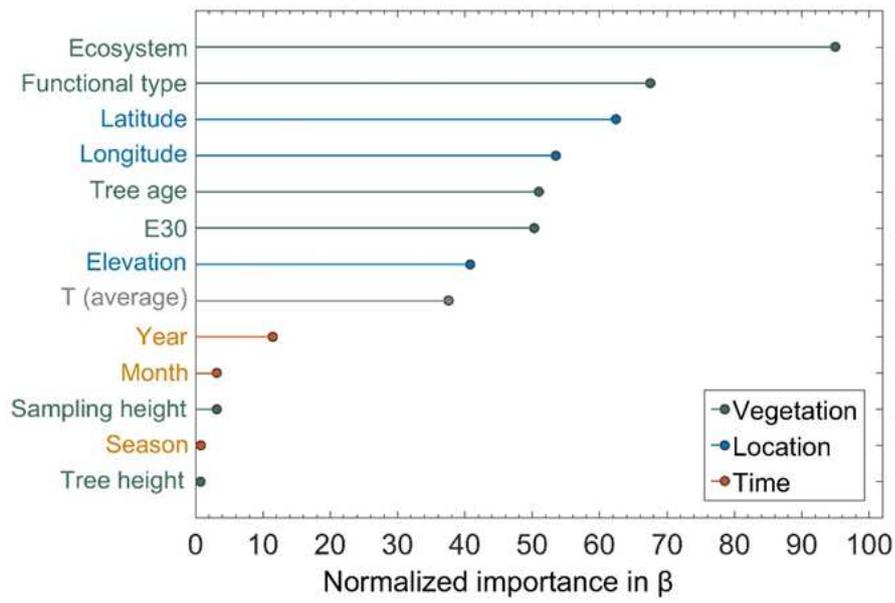
577



578

579

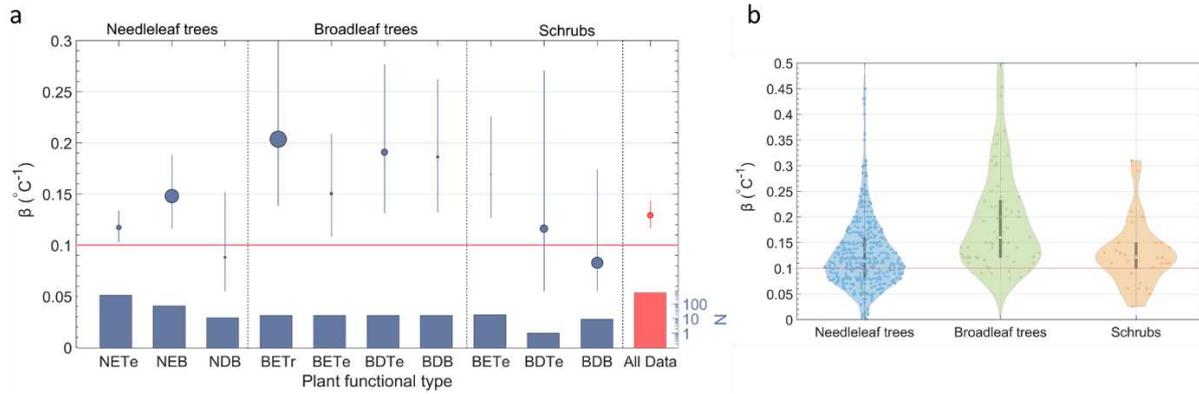
580 **Figure 1.** Groups of coefficients of determination (R^2) for experimentally derived
581 dependencies (β) of MT emissions on temperature. The blue and green error bars indicate
582 95% confidence intervals. On the right, the distribution of the N=696 values of β extracted
583 from the literature is displayed together with a boxplot that illustrates their median (white
584 circle) and 25th and 75th percentiles (lower and upper bounds of the box), respectively. Data
585 below the red dashed line (on the right) were obtained from regressions with poor goodness of
586 fit ($R^2 < 0.2$) and were excluded from further analysis.



587

588

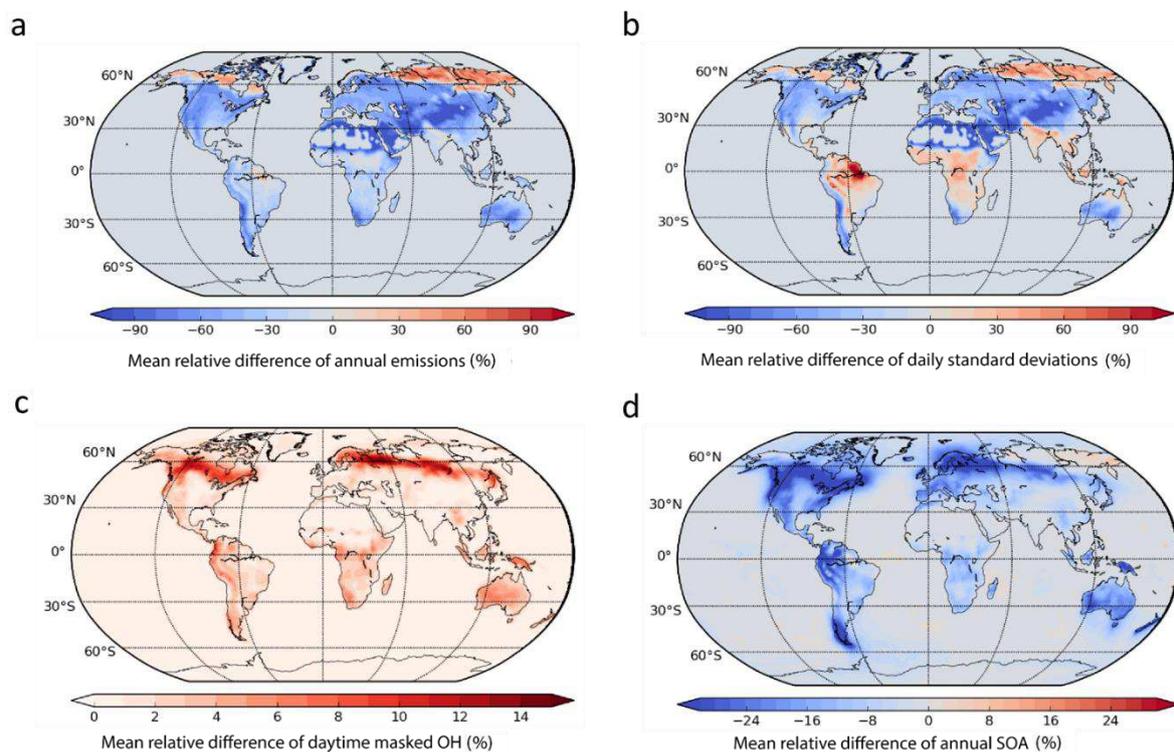
589 **Figure 2.** Feature importance ranking. Indicative contribution of individual parameters of
 590 vegetation, location, and time parameters to the variation of the β coefficient.



591

592

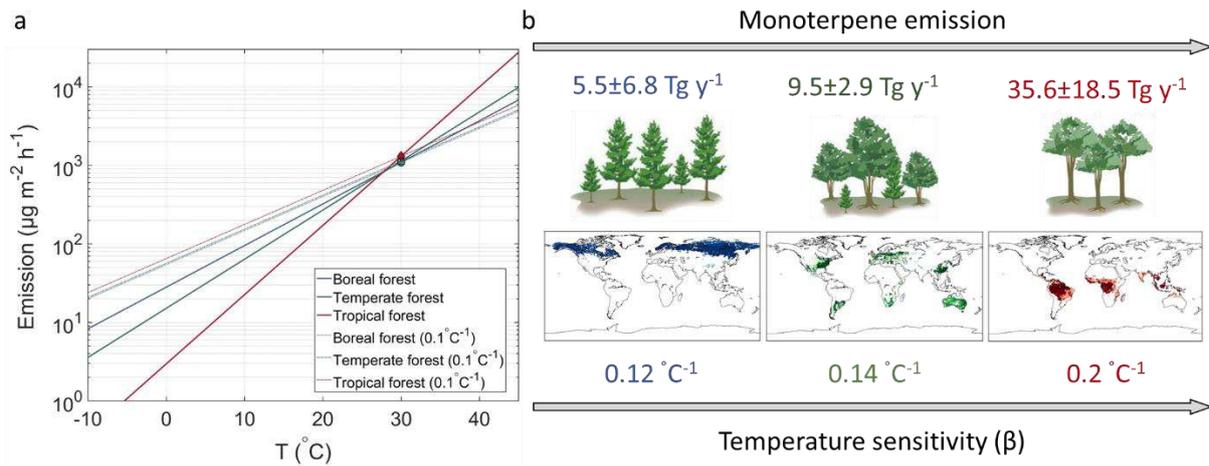
593 **Figure 3.** Plant functional type-dependent β coefficients. Dependence on temperature (β) for
 594 the plant functional types (A) and wider categories of woody plants (B). The error bars in
 595 panel a indicate 95% confidence intervals, and the size of the blue bullet points is proportional
 596 to the global surface area of each plant functional type. NETe, needleleaf evergreen temperate
 597 forest; NEB, needleleaf evergreen boreal forest; NDB, needleleaf deciduous boreal forest;
 598 BETr, broadleaf evergreen tropical forest; BETe, broadleaf evergreen temperate forest/shrubs;
 599 BDTe, broadleaf deciduous temperate forest/shrubs; BDB, broadleaf deciduous boreal
 600 forest/shrubs.



601

602

603 **Figure 4.** Model simulations. Annual mean relative differences (β -factor = $0.1 \text{ } ^\circ\text{C}^{-1}$ and β -
 604 factor from Fig. 3 based on all data) derived from hourly data over a climatic year. A. Total
 605 monoterpene emissions, B. Standard deviation of daily emissions, C. Daytime OH, and D.
 606 SOA production.



607

608 **Figure 5.** Monoterpene emissions over temperature gradients for the three main ecosystems.
 609 A. Emissions using the standardized potential emission at 30 °C and β in MEGAN v2.1
 610 (dashed lines) and using β derived from this study (panel b and Supplementary Table 2). The
 611 maps in panel B represent the global coverage of boreal (blue), temperate (green), and tropical
 612 (red) forests.

Supplementary Files

This is a list of supplementary files associated with this preprint. Click to download.

- [Bourtsoukidisetal.Supplement.pdf](#)
- [Bourtsoukidisetal.Supplement.pdf](#)