

Identification of volatile organic compounds and their sources driving ozone and secondary organic aerosol formation in NE Spain

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1 Abstract

2 Volatile organic compounds (VOCs) play a crucial role in the formation of ozone (O₃) and secondary
3 organic aerosol (SOA). In NE Spain, elevated O₃ levels have been observed and attributed to factors
4 such as low summer precipitation, high insolation, elevated levels of NO₂ and anthropogenic VOC
5 pollution, substantial biogenic VOC emissions, and the vertical recirculation of air masses influenced
6 by the combination of orographic features and meteorological patterns. We conducted
7 measurements of VOC ambient mixing ratios during both summer and winter at two stations: a
8 Barcelona urban background station (BCN) and the Montseny rural background station (MSY).
9 Subsequently, we employed positive matrix factorization (PMF) to analyze the VOC mixing ratios and
10 identify their sources. Our analysis revealed five common sources: anthropogenic I (traffic &
11 industries); anthropogenic II (traffic & biomass burning); isoprene oxidation; monoterpenes; long-
12 lifetime VOCs. To assess the impact of these VOCs on the formation of secondary pollutants, we
13 calculated the ozone formation potential (OFP) and secondary organic aerosol formation potential
14 (SOAP) associated with each VOC. In conclusion, our study provides insights into the sources of VOCs
15 and their contributions to the formation of ozone and SOA in NE Spain. The OFP was primarily
16 influenced by anthropogenic aromatic compounds from the traffic & industries source at BCN (38–
17 49%) and during winter at MSY (34%). In contrast, the summer OFP at MSY was primarily driven by
18 biogenic contributions from monoterpenes and isoprene oxidation products (45%). Acetaldehyde (10–
19 35%) and methanol (13–14%) also made significant OFP contributions at both stations. Regarding SOA
20 formation, anthropogenic aromatic compounds originating from traffic, industries, and biomass
21 burning played a dominant role (88–93%). However, during summertime at MSY, monoterpenes
22 became the primary driver for SOA formation (41%). These findings emphasize the importance of
23 considering both anthropogenic and biogenic VOCs in air quality management strategies.

24 1.0 Introduction

25 Volatile organic compounds (VOCs) play an important role in the formation of ozone (O₃) and
26 secondary organic aerosol (SOA). O₃ is formed in the troposphere by photochemical reactions
27 involving the oxidation of VOCs, including methane (CH₄) and carbon monoxide (CO) (Jacob, 1999a,
28 1999b). It has been estimated that approximately 90% of global tropospheric O₃ is formed by the
29 oxidation of VOCs (Jacob, 1999b, 1999a; Möller, 2004; Stevenson et al., 2006; Young et al., 2013).
30 Meanwhile, SOAs are primarily formed via the gas-to-particle conversion mechanism from precursor
31 gases, which include VOCs, NH₃, H₂SO₄, HNO₃, and HCl, among others (Jacob, 1999c; Seinfeld and
32 Pandis, 2016).

33 After the success of the EU policy for abating air pollutants, which caused a reduction in the ambient
34 air concentrations of several (especially primary) pollutants, interest in secondary air pollutants has
35 increased in both the science and policy realms. Notably, the ambient concentrations of secondary
36 pollutants, such as O₃ and SOA, did not follow a decreasing trend (EEA, 2019, 2018a; EMEP/CCC, 2016;
37 in 't Veld et al., 2021; Paoletti et al., 2014; Querol et al., 2016, 2014), which is detrimental to human
38 health. The European Environmental Agency (EEA, 2021) reported 307,000 premature deaths in EU-
39 28 countries attributable to fine particulate matter (PM_{2.5}) in 2019, 40,400 attributable to NO₂, and
40 16,800 attributable to O₃. Furthermore, O₃ also has a considerable negative effect on vegetation
41 (Felzer et al., 2007; Krupa and Manning, 1988). It has been estimated that by 2030, there will be a
42 global loss of €14–29 billion annually due to crop losses caused by O₃ (Avnery et al., 2011).

43 This is especially worrying for the western Mediterranean basin, which is prone to elevated O₃ levels
44 due to low summer precipitation, high insolation, high NO₂ pollution, high biogenic VOC emissions,
45 and the vertical recirculation of air masses caused by the combination of orographic features and
46 meteorological patterns (Gangoiti et al., 2001; Millán, 2014; Millán et al., 2002, 1997; Pérez et al.,
47 2004; Querol et al., 2017). Querol et al. (2016) reported that between 2000 and 2015, an O₃ level
48 increase of 1.2% yr⁻¹ occurred at industrial and urban background sites in Spain, with an increase of

49 1.9% yr⁻¹ at traffic sites. Rural background concentrations did not show a significant trend but were
50 higher in comparison to the other sites (urban, traffic, and industrial). Past studies have shown a
51 relative increase in SOA throughout the study area. For example, Via et al. (2021) compared organic
52 aerosol measurements using an aerosol chemical speciation monitor (ACSM) between May 2014 and
53 May 2015, as well as between Sep 2017 and Oct 2018, showing an increase in the relative SOA content
54 in PM₁ and a higher degree of oxidation during the latter period. The same observation was made by
55 in 't Veld et al. (2021), who estimated relative SOA increases of 12 and 8% in the chemical composition
56 of PM_{2.5} in urban and rural background stations, respectively, between 2009 and 2018. Furthermore,
57 secondary pollutants are closely linked to one another since higher O₃ levels increase the readily
58 available hydroxyl- and nitrate-oxidizing radicals (OH[•] and NO₃[•]) in the urban atmosphere. An example
59 of this was observed in downtown Madrid, where a 30–40% increase in O₃ was measured alongside
60 OH[•] and NO₃[•] increases of up to 70 and 90%, respectively, between 2007 and 2014 (Saiz-Lopez et al.,
61 2017). Additionally, a reduction in sulfur dioxide (SO₂) (EEA, 2015) and nitrogen oxides (NO_x) (EEA,
62 2018b) in Europe between 1990 and 2011 also led to a decrease in the consumption of the oxidizing
63 radicals, which increased their availability and might have resulted in an atmosphere more prone to
64 SOA production (Querol et al., 2018; Saiz-Lopez et al., 2017).

65 Increased knowledge regarding the relationships of VOC sources and their potential to form O₃ and
66 SOA is required to develop cost-effective abatement strategies in the western Mediterranean. VOCs
67 encompass a wide variety of organic compounds, which include over 1,000 chemical components
68 originating from a variety of sources. Although previous studies have investigated the VOCs occurring
69 in NE Spain (Filella and Peñuelas, 2006; Seco et al., 2013, 2011; Yáñez-Serrano et al., 2021a) and
70 contributed to a greater understanding of the VOCs in this region, source apportionment was only
71 performed in a rural background site during summer (Yáñez-Serrano et al., 2021a), while a comparison
72 between an urban background and rural background was only performed during winter (Seco et al.,
73 2013). Finally, most published studies have not investigated the potential of existing VOCs to form O₃
74 or SOA.

75 This study focuses on evaluating the levels and source contributions of VOCs in NE Spain by using
76 proton-transfer-reaction mass spectrometry (PTR-MS) to measure online VOC mixing ratios. This
77 allowed us to identify and quantify a selection of VOCs in both the Barcelona urban background site
78 and the Montseny rural background site. The differences between the stations allowed us to highlight
79 differences between an urban and rural environment and show which VOCs are emitted locally and
80 which are emitted regionally. Furthermore, the extensive measurements conducted during both the
81 summer and winter months at both stations allowed us to identify seasonal patterns. The use of both
82 the O₃ formation potential (OFP) and the SOA formation potential (SOAP) to estimate the amount of
83 O₃ and SOAs that VOCs can form under ideal situations, respectively, permits the identification of the
84 most potentially important precursors to both O₃ and SOA formation in the study area. Additionally,
85 to aid in the identification of VOC sources, a positive matrix factorization (PMF) model was applied on
86 the VOC dataset to obtain the sources.

87 2.0 Methodology

88 2.1 Sampling location

89 VOC measurements were performed using PTR-MS at two stations in the conurbation of Barcelona in
90 NE Spain. Both sites are included in the European Aerosols, Clouds, and Trace Gases Research
91 Infrastructure Network (ACTRIS) and the Air Quality Monitoring Network (AQMN) of the Catalan
92 Government, while MSY is part of the Global Atmosphere Watch (GAW) network of the World
93 Meteorological Organization (WMO) (Figure 1). A mountain in the Montseny Natural Park, located in
94 El Vilar de la Castanya (el Brull), is home to one of the two stations, which is a rural background station
95 (hereafter MSY; 41°46'45.63"N, 02°21'28.92"E; 720 m a.s.l.). The MSY station is located 40 km from
96 the Mediterranean coast and 50 km to the N-NE of Barcelona. It has been demonstrated to be
97 representative of the regional ambient background and is sufficiently elevated and distant from
98 specific urban anthropogenic emission sources (Cusack et al., 2012; in 't Veld et al., 2023, 2021;
99 Pandolfi et al., 2016, 2014; Pérez et al., 2008; Pey et al., 2009; Ripoll et al., 2015), although the station

100 may be impacted by emissions from urban and industrial areas during anticyclonic atmospheric
101 conditions due to wind circulation (Ealo et al., 2018; Gangoiti et al., 2001; Millán et al., 2002, 1997;
102 Pandolfi et al., 2016, 2013). The second station is an urban background air quality station (hereafter
103 BCN; 41°23'14.5"N 2°06'55.6"E; 68 m asl) located at the Institute of Environmental Assessment and
104 Water Research (IDAEA-CSIC) in Barcelona. It is positioned adjacent to Diagonal Avenue, one of the
105 city's main roads, and is flanked by a variety of commercial activities (Amato et al., 2009; Cusack et al.,
106 2012; in 't Veld et al., 2023, 2021; Pandolfi et al., 2016, 2014; Pérez et al., 2008; Pey et al., 2009; Querol
107 et al., 2014, 2004a, 2004b; Ripoll et al., 2015). At both measurement sites, an aethalometer (Magee
108 Scientific, USA, model AE33) was used to quantify the aerosol light absorption coefficient at seven
109 different wavelengths (370, 470, 520, 590, 660, 880, and 950 nm) and the mass concentration of black
110 carbon (BC) (Drinovec et al., 2015). Additionally, the AQMN used instrumentation to measure hourly
111 concentrations of O₃, CO, NO, NO₂, and NO_x (Table S6). Finally, ambient temperature and solar
112 radiation were measured at the station in MSY, and at the Department of Physics building of the
113 Barcelona University for the BCN station. Wind speed and direction data were also usually included;
114 however, this was unavailable at MSY during the 2022 measurement period and is thus not discussed
115 in this paper.

116 2.2 PTR-MS sampling

117 The VOC mixing ratios in ambient air were measured using PTR-MS. Different instruments were used
118 depending on the location and year. In 2022, a time-of-flight mass spectrometer (PTR-ToF-MS, Ionicon
119 PTR TOF 4000X2, Ionicon Analytik, Innsbruck, Austria) was used in Barcelona, while a PTR-MS with a
120 quadrupole mass spectrometer (PTR-MS, Ionicon PTR-MS Quad, Ionicon Analytik, Innsbruck, Austria)
121 was used at MSY. The measurements at MSY in 2017 used a PTR-ToF-MS (Ionicon PTR TOF 8000,
122 Ionicon Analytik, Innsbruck, Austria). A detailed description of the instrument is provided by Graus et
123 al. (2010). In summary, proton transfer is a form of soft chemical ionization from hydronium ions
124 (H₃O⁺), with the distinct feature that H₃O⁺ ionizes compounds with little energy excess, resulting in
125 minor fragmentation, thus making it an excellent method to measure ambient air (Dunne et al., 2018;

126 Gouw et al., 2003; Ionicon Analytik GmbH, 2014; Warneke et al., 2011). The parameters of each PTR-
127 MS are described in Table 1. For the 2022 measurements, a periodic calibration was made during the
128 measurements, with automatic blanks taken every hour for MSY and every 6 hours for BCN. For the
129 measurement in 2017, a calibration was performed at the start of the measurement. Since a blank
130 measurement was missing for the measurement period, an approximation of the blank was made
131 using blanks from the previous campaign using the same apparatus. Table 1 presents the
132 measurement periods at both stations, the apparatus used, and the measurement frequency. The
133 periods were selected as approximately 1 month in duration when data were available at both BCN
134 and MSY.

135 Two gravimetrically prepared multicomponent standards created by Apel-Riemer (Apel-Riemer
136 Environmental Inc., Miami, USA) and Restek (Restek Corporation, Bellefonte, PA, USA) contained all
137 VOC compounds discussed here, which were humidity-dependent calibrated at various dilution steps.
138 These calibrations used bubbled zero air to dilute the standards and were regulated as closely as
139 possible to the ambient humidity. A list of the compounds is presented in Table 2. Notably, since PTR-
140 MS measures the exact masses of compounds, it is not compound-specific since compounds with the
141 same masses may interfere with the measured signal. The compounds were assigned to the masses
142 based on their exact mass, the compounds included in the calibration canister, the GLOVOCs database
143 (Yáñez-Serrano et al., 2021b), and previous measurements in the study area (Peñuelas et al., 2009;
144 Seco et al., 2013, 2011; Yáñez-Serrano et al., 2021a). Nevertheless, some other compounds might
145 contribute to these masses during ambient measurement. While most masses had a single compound
146 assigned to them, a few masses had multiple compounds assigned to them. A mass-to-charge ratio
147 (m/z) of 59 is the same for both acetone and propanal. However, the ambient concentrations of
148 propanal are significantly lower when compared to acetone (Hellén et al., 2004; Mingxiao Li et al.,
149 2018). Therefore, in this study, most instances of m/z 59 will be assumed to be acetone. A similar
150 situation occurred with m/z 69, which was primarily attributed to isoprene but has possible
151 interference from furan and fragments of 2-methyl-3-butene-2-ol (MBO); however, these

152 concentrations are much lower when compared to isoprene. Therefore, m/z 69 will be assumed to be
153 isoprene in this study (Juráň et al., 2017; Karl et al., 2012; Kaser et al., 2013). Meanwhile, m/z 71 can
154 be both methyl vinyl ketone (MVK) and methacrolein (MACR) since both are isomers of C₄H₆O.
155 Therefore, this mass was considered to be a combination of these two VOCs. Moreover, m/z 107 can
156 be either ethylbenzene, *o*-, *p*-, or *m*-xylene, and is thus referred to as C₈ aromatics since they are all
157 isomers of C₈H₁₀. Another common VOC with m/z 107 is benzaldehyde (C₇H₆O). While benzaldehyde
158 has a different exact mass and can be distinguished using PTR-ToF-MS, it cannot be ruled out that it
159 interferes with the C₈ aromatics peak when using quadrupole-based PTR-MS instruments. Finally, m/z
160 137 contains various monoterpenes that are isomers of C₁₀H₁₆ and are thus referred to as
161 monoterpenes. The uncertainty of the measurements was calculated according to the error
162 propagation approach (Doerffel, 1984). This method calculates the total uncertainty of the PTR-MS
163 while considering the calibration uncertainty (which includes multicomponent gas standard and mass
164 flow controller errors), the PTR-MS instrument, and the background error. Table 2 also reports the
165 median uncertainty in % based on the uncertainty from the calibration gas, blanks, and measurements
166 combined.

167 2.3 PMF source apportionment

168 The source apportionment analysis was performed on the VOC dataset via PMF using the U.S.
169 Environmental Protection Agency's PMF v5.0 software (Norris et al., 2014). PMF is a multivariate
170 factor analysis tool that decomposes time trends in chemical composition into factor contributions
171 and chemical profiles by performing chemical mass balance between measured species
172 concentrations and the sum of source contributions for those species:

$$173 \quad X_{ij} = \sum_{k=1}^p g_{ik} * f_{kj} + e_{ij}$$

174 Here, X is the data matrix with i number of samples and j number of chemical species. p is the number
175 of sources, f is the chemical profile of each source with mass contribution g , and e_{ij} is the residual for
176 each sample. To obtain better factor profiles, certain VOCs were set as “weak” species, which means
177 that the uncertainty is tripled. These species were selected based on the signal-to-noise ratio ($0.5 >$
178 $S/N > 1.5$) and visual anomalies in the factor profiles. Strong species included acetone, isoprene,
179 benzene, toluene, and monoterpenes. Weak species included methanol, acetonitrile, acetaldehyde,
180 MVK+MACR, methyl ethyl ketone (MEK), and C₈ aromatics.

181 In this study, a multisite solution was used by aggregating the BCN and MSY data into a single dataset.
182 The main benefits of running a multisite PMF were the inclusion of a larger dataset when compared
183 to separate single-site PMF models. A multisite PMF produces more robust results and can obtain
184 source profiles common to both sites, thus allowing a direct comparison between the two stations.
185 This has not been previously done for VOC measurements in the study area but has been successfully
186 applied to the chemical speciation of particulate matter (in 't Veld et al., 2023, 2021; Pandolfi et al.,
187 2016). However, we also acknowledge the limitations of this analysis stemming from the physical
188 distance between the two stations, possible differences in local source types and their magnitude, and
189 different chemical profiles for the same types of sources (Escrig et al., 2009). The results were
190 bootstrapped 100 times, and a displacement model was used to determine the uncertainties in the
191 PMF results.

192 2.4 Ozone formation potential

193 Ozone formation potential (OFP) has frequently been used to assess the maximal possible contribution
194 to photochemical ozone formation from particular VOCs under ideal reaction circumstances (Carter,
195 2010, 2009, 1994; Venecek et al., 2018). This is known as the maximum incremental reactivity (MIR).
196 The MIR values in this study were obtained from Venecek et al. (2018), which was based on the
197 California statewide air pollution research center (SAPRC) box model that uses a representative urban
198 atmosphere (Carter, 1994; Venecek et al., 2018).

199
$$OFP_i = C_i * MIR_i$$

200 Here, the OFP of VOC *i* is calculated by multiplying the ambient concentration C_i with the MIR of VOC
201 *i*.

202 2.5 Secondary organic aerosol potential

203 The SOAP was calculated in a similar manner, where the SOAP of VOC *i* was calculated by multiplying
204 the ambient concentration C_i with the SOA yield $Y_{SOA,i}$.

205
$$SOAP_i = C_i * Y_{SOA,i}$$

206 The SOA yields were obtained from Gu et al. (2021), who experimentally determined the SOA yields
207 from smog chamber experiments for a broad inventory of VOCs based on the two-product model (Gu
208 et al., 2021; Odum et al., 1996; Wu and Xie, 2018). In the case of no VOC data being available, they
209 obtained the SOA yields from Wu et al. (2017) or used the SOA yield of a VOC with a similar structure.
210 The SOAP was converted to μgm^{-3} in order to obtain the SI units often used for particle concentrations.

211
$$SOAP_{\mu\text{gm}^{-3}} = \frac{P * MW * SOAP_{\text{ppb}}}{62.4 * (273.2 + T)}$$

212 Here, *P* is the atmospheric pressure in mmHg, which was set at 706 mmHg (equivalent to 1 atm). *MW*
213 is the molecular weight of the VOC, while $SOAP_{\text{ppb}}$ is the SOAP calculated in ppb. This was divided by
214 the ideal gas constant of 62.4 torr/mol, multiplied by the ambient temperature in K.

215

216 2.6 Correlations

217 Statistical analysis was performed using the SciPy package for Python (Virtanen et al., 2020). Using this
218 package, Pearson's correlation coefficient (*r*) was calculated between the VOCs, gases, BC, PMF
219 sources, and meteorological data (Figure S1-4). The statistical significance of each correlation was also
220 calculated and denoted accordingly when $p < 0.05$.

221 3.0 Results and Discussion

222 3.1 Regulated gaseous pollutants

223 The gas data from the two stations showed that anthropogenic gaseous pollutants, such as NO₂ and
224 CO, were much higher at BCN when compared to MSY (Figure 2). At BCN, NO₂ levels had two distinct
225 peaks. One peak was in the morning, and another was during the evening rush hour, which was also
226 observed for CO. Furthermore, the concentrations during weekends were much lower when
227 compared to the weekly average due to the lower traffic levels during weekends (in 't Veld et al., 2021;
228 Massagué et al., 2019). At MSY, peaks during working days were missing. Nonetheless, NO₂ showed
229 an increase in the evening that is considerably lower during the weekend, which suggests the
230 potential entrance of pollutants from nearby areas (in 't Veld et al., 2023, 2021; Pandolfi et al., 2016;
231 Querol et al., 2013; Seco et al., 2011; Yáñez-Serrano et al., 2021a). The CO concentrations did not show
232 a diel cycle due to the low accuracy of our measurements in combination with the low variability in
233 CO concentrations; however, they did show higher levels during winter when compared to summer.
234 This seasonal trend was also observed at BCN due to the increase in anthropogenic emissions during
235 winter, such as the increased use of residential heating (Reche et al., 2012; Viana et al., 2013) and a
236 higher boundary layer height (Pérez et al., 2008).

237 The O₃ diel cycle showed a clear photochemical pattern at both stations with increasing levels during
238 the day, which peaks around midday, followed by a decrease (Figure 3). Tropospheric O₃ is formed in
239 the atmosphere due to the photolysis of NO₂ to NO and O[•]. In turn, the O[•] can react with O₂ to form
240 O₃ (Jacob, 1999b, 1999a; Möller, 2004; Stevenson et al., 2006; Young et al., 2013). Since this process
241 requires sunlight, it only occurs during daylight, showing an increase during daytime and a decrease
242 after the midday peak (Figure 3). Furthermore, NO₂ concentrations also increase during the day at
243 BCN due to the oxidation of NO by O₃, as well as anthropogenic emissions (Gligorovski et al., 2015;
244 Jacob, 1999c, 1999b; Kim et al., 2013; Levy, 1971; Mengze Li et al., 2018; McElroy, 2002). As expected,
245 O₃ concentrations were higher during the summer due to higher temperatures and stronger solar

246 radiation (Monks et al., 2015), but also to the specific atmospheric dynamics of the western
247 Mediterranean during this season, resulting in the vertical recirculation of air masses along the coast
248 and throughout the basins (Millán et al., 2002, 1997). O₃ levels were also generally higher at MSY (76.7
249 ± 23.5 ppbv) when compared to BCN (58.4 ± 25.7 µgm⁻³). MSY typically has higher levels of O₃ than
250 the urban site of BCN due to frequent urban pollution plume transport from BCN to MSY and the less
251 effective NO titration of O₃ at MSY when compared to BCN (Massagué et al., 2019; Querol et al., 2016).

252 3.2 VOC mixing ratios

253 Table 3 presents the 11 compounds, their masses, and their average (taken over the whole campaign,
254 Section 1 of the SI) mixing ratios separated by season and station. In all cases (both at BCN and MSY
255 during both seasons), methanol was the most abundant VOC, which mostly originated from biogenic
256 emissions and secondary formation (Holzinger et al., 2005, 1999; Lewis et al., 2005; Seco et al., 2007;
257 Sjostedt et al., 2012). Acetone was the second-most abundant VOC and, like methanol, mostly
258 originated from biogenic emissions and secondary formation (Arnold et al., 2004; Holzinger et al.,
259 2005, 1999; Lewis et al., 2005; Seco et al., 2007; Singh et al., 1994; Sjostedt et al., 2012). The high
260 levels of these short-chain oxygenated VOCs can be attributed to their low reactivity with atmospheric
261 oxidants, especially when compared to the other set of VOCs (Table 6) (Arnold et al., 2004; Filella and
262 Peñuelas, 2006; Lewis et al., 2005; Seco et al., 2013), which will be discussed in detail in subsequent
263 sections.

264 The average mixing ratios of this study were compared to previous studies conducted in the study
265 area between 2009 and 2019 (Section 2 of the SI). The comparison between these studies showed
266 that the mixing ratios at MSY were stable between 2009 and 2019, indicating consistency in the
267 regional air quality regarding VOCs. At BCN, differences were observed, where the different
268 measurement periods and protocols used could have contributed to this. Additionally, since BCN was
269 located closer to the sources in comparison to MSY, the measurements might be more prone to the

270 variations of local sources from one campaign to another, especially regarding the mean values of
271 short-period measurement campaigns.

272 In summer, all VOCs at MSY show a small peak in the evening (Figure 3). This is a currently unknown
273 anomaly of the MSY station, which was also observed in the temperature and radiation data. This
274 phenomenon was also observed by Yáñez-Serrano et al. (2021a), who sampled VOCs a few kilometers
275 from the measurement station, which indicates that this is not an anomaly of the measurement
276 station.

277 3.3 Source apportionment

278 To determine the sources of all VOCs, we performed a multisite PMF source apportionment. The
279 model was applied to an aggregated dataset from both the summer and winter periods of both
280 stations. This made it possible to compare source profiles between the different seasons and stations.
281 The source apportionment analysis identified five factors (Table 4), with the profiles presented in
282 Figure S5. For ease of interpretation, all factor names are written in *italics* in this paper. The optimal
283 number of sources was selected by inspecting the Q values, residuals, G space plots, and physical
284 meanings of the factors. To confirm the optimal factor profiles of the PMF model, the data were also
285 bootstrapped. Each dataset was bootstrapped 100 times with a minimum correlation R-value of 0.6.
286 Finally, the model error was estimated using the base model displacement method.

- 287 • *Factor 1: Anthropogenic I: traffic & industry factor*
 - 288 ○ This source originated from traffic and industry emissions at BCN and was mostly long-
289 distance transport at MSY.
- 290 • *Factor 2: Anthropogenic II: traffic & biomass burning factor*
 - 291 ○ Additional anthropogenic source that accounted for a second traffic source and
292 biomass burning.
- 293 • *Factor 3: Isoprene oxidation factor*

294 ○ Contains isoprene and its oxidation products, which are both anthropogenic and
295 biogenic at BCN, and mostly biogenic at MSY.

296 • *Factor 4: Monoterpene factor*

297 ○ A biogenic source at both stations that contained only monoterpenes.

298 • *Factor 5: Long-lifetime VOC factor*

299 ○ A source combining the VOCs with long atmospheric lifetimes. These were mostly
300 anthropogenic at BCN, while biogenic and formed due to photo-oxidation at MSY. This
301 source also contained VOCs that could not be attributed to the other four factors.

302 To confirm the findings of the multisite PMF, separate PMF analyses were performed for both the BCN
303 and MSY stations, combining both seasons. All sources were identified in separate PMFs with the
304 following notes. At MSY, the biogenic sources were identified as a single source combining isoprene,
305 MVK+MACR, and monoterpenes. At BCN, the biogenic sources were separated into two different
306 sources, as observed in the multisite solution. As will be described later in this section, this difference
307 is due to the following factors. In the BCN area, isoprene can be emitted from biogenic and
308 anthropogenic sources; however, such an anthropogenic source was negligible at MSY. Furthermore,
309 the *long-lifetime VOC* factor was separated into two distinct groups at MSY, with one being
310 characterized by photo-oxidized VOCs, which was lacking in the BCN PMF—and thus in the multisite
311 PMF. All other sources were identified at both stations with the same tracers.

312

313

314 3.3.1 Factor 1: Anthropogenic I: Traffic & Industry

315 *Factor 1* was mostly traced by toluene (81% of the total toluene) and C₈ aromatics (67%), with tracers
316 of benzene (25%) and isoprene (19%, discussed in more detail in *Factor 3*). These VOCs are mostly
317 emitted anthropogenic VOCs. The high content of aromatic hydrocarbons and oxygenated compounds
318 without biogenic VOCs further indicates the anthropogenic origin of this factor.

319 The diel cycle of BCN (Figure 4) showed two clear traffic peaks during both seasons (8:00 and 20:00),
320 coinciding with those of CO and NO₂ (Figure 2). Previous studies at BCN determined that
321 anthropogenic VOC emissions are primarily caused by road traffic (Filella and Peñuelas, 2006; in 't Veld
322 et al., 2021; Jiménez et al., 2005). The lower concentrations during weekends further confirm this
323 since anthropogenic emissions from traffic and industry are significantly lower during weekends when
324 compared to weekdays. This phenomenon was also observed for CO and NO₂ (Filella and Peñuelas,
325 2006; in 't Veld et al., 2021). Moreover, these phenomena were also observed for the major
326 constituents of this factor, toluene and C₈ aromatics (**Error! Reference source not found.**Figure 3),
327 indicating that they both drive this factor.

328 These observations point to the anthropogenic origin of this factor, which is further confirmed by the
329 toluene to benzene (T/B) ratio. Generally, high T/B ratios are observed in industrial areas (Seco et al.,
330 2013; Tan et al., 2021), while lower ratios would be observed as fresh vehicle emissions increase (Heeb
331 et al., 2000; Khoder, 2007; Langford et al., 2009; Pinthong et al., 2022). If the T/B ratio fell below 1,
332 the main source of emissions would come from biomass burning (Liu et al., 2008; Tan et al., 2021). In
333 both stations, the T/B ratio was lower in the winter when compared to the summer (Figure 5). At BCN,
334 the T/B ratio was 9.73 ± 5.16 during summer, which indicates that this factor is dominated by industrial
335 emissions mixed with traffic emissions since industrial emissions are characterized by low levels of
336 benzene and high levels of toluene (Filella and Peñuelas, 2006; Seco et al., 2013). During winter, the
337 T/B ratio was lower with 2.92 ± 2.28 , which is in line with previous studies conducted in the study area
338 (Filella and Peñuelas, 2006; Seco et al., 2013). The lower values during winter were also due to an

339 increase in benzene emissions from residential heating, which will be discussed further in *Factor 2*.
340 During both seasons, a decrease in the T/B ratio could be observed due to the substantial decrease in
341 traffic emissions during the weekend (Jiménez et al., 2005). The decrease between summer and winter
342 was also observed at MSY, where the T/B ratio was 3.58 ± 3.50 during summer compared to $2.44 \pm$
343 2.91 in winter due to the same phenomena. The lower values at MSY compared to BCN were a result
344 of photochemical aging since toluene is scavenged more rapidly than benzene (Table 6) (Gelencsér et
345 al., 1997; Seco et al., 2013) since MSY is located further from the sources of traffic and industry, which
346 will most likely involve transported air masses and result in a lower T/B ratio.

347 Notably, anthropogenic emissions are generally higher during winter than summer at BCN, as
348 observed for NO₂ and CO (Figure 2) (in 't Veld et al., 2021; Pandolfi et al., 2016; Pérez et al., 2016;
349 Querol et al., 2014, 2001; Viana et al., 2013) and previous studies of benzene and toluene at BCN
350 (Filella and Peñuelas, 2006). Contrary to this, in our study, the mixing ratios of toluene and the C₈
351 aromatics in the atmosphere were higher during summer at BCN, while the winter levels were higher
352 at MSY (Table 3). Currently, it remains unknown why this increase occurred during summer.

353 *Factor 1* showed a high correlation with other anthropogenic contaminants in the study area, such as
354 BC concentrations (summer: $r = 0.73$; winter: $r = 0.83$), NO₂ concentrations (summer: $r = 0.46$; winter:
355 $r = 0.66$), and other anthropogenic factors, namely the *traffic & biomass burning factor (Factor 2)*
356 (summer: $r = 0.55$; winter: $r = 0.89$) and the *isoprene oxidation factor (Factor 3)* (summer: $r = 0.42$;
357 winter: $r = 0.72$), which will be discussed in their respective sections. Overall, *Factor 1* represents
358 emissions from traffic and industry at BCN.

359 At MSY, the mixing ratios of *Factor 1* were considerably lower when compared to BCN due to its
360 distance from anthropogenic sources. During summer, levels were negligible and increased in winter
361 due to greater anthropogenic emissions. During summer at MSY, the *traffic & industry factor* showed
362 some correlation with solar radiation ($r = 0.42$), with increasing concentrations during daylight hours
363 when compared to the evening (Figure 5), while the two peaks that coincided with rush hours at BCN

364 were absent. All other correlations with anthropogenic contaminants were not statistically significant.
365 This indicates that instead of an anthropogenic origin, the factor might have a minor biogenic
366 contribution to MSY. This should be considered an option since toluene could be emitted by
367 Mediterranean holm oaks and pines (Heiden et al., 2009; Holzinger et al., 2000; Misztal et al., 2015).
368 Figure S6 presents the diel cycle of toluene and temperature for each day. Each diel cycle shows a
369 peak during midday on most measured days, which was absent for BCN. This correlation was only
370 present for toluene and missing for the C₈ aromatics. Although the diel cycle showed a similar pattern
371 for both the toluene and the C₈ aromatics, the correlation between these two compounds at MSY was
372 not as high ($r = 0.52$, compared to $r = 0.94$ at BCN for the same season). Both VOCs still had a significant
373 correlation with other anthropogenic tracers (Figure S3-4), which would indicate an anthropogenic
374 origin for this factor with a minor biogenic contribution from toluene. Yáñez-Serrano et al. (2021)
375 identified a similar factor at a site located nearby the MSY station and attributed it to traffic emissions,
376 which was the lowest contributing factor during summer, as we observed in our study. Tan et al. (2021)
377 also found an anthropogenic source in Hong Kong with the same tracers, attributing it to industrial
378 and traffic emissions.

379 3.3.2 Factor 2: Anthropogenic II: Traffic & Biomass Burning

380 *Factor 2* was traced by benzene (69%), acetaldehyde (67%), and acetonitrile (36%), as well as various
381 oxidized VOCs (OVOCs), such as MEK (27%), methanol (19%), and C₈ aromatics (19%). Both benzene
382 and acetonitrile are typical markers of biomass burning, with studies also showing that acetaldehyde,
383 MEK, and methanol can originate from biomass burning (Cerqueira et al., 2013; Holzinger et al., 1999;
384 Li et al., 2014; Tan et al., 2021).

385 The mixing ratio of this factor was much higher during winter when compared to summer at both
386 stations. While biomass burning can have various sources, such as forest fires and the burning of
387 agricultural waste, the increase during winter indicated that this increase was due to residential
388 heating, which was previously detected in the study area (Reche et al., 2012; Viana et al., 2013). The

389 T/B ratio also confirmed this, with lower ratios during the winter season when compared to the
390 summer at both stations, indicating a higher influence of biomass burning (Figure 7).

391 Figure 7 presents a scatterplot of the benzene and toluene mixing ratios. While they are mostly
392 grouped together, showing a high T/B ratio, two periods deviated from this (marked in red). The first
393 was at BCN on the 14th of February 2022 between 07:00 and 10:00 and the second was at MSY on the
394 13th of June 2022 between 17:00 and 21:30. During these events, the T/B ratio differentiated from the
395 group to a ratio below 1, indicating biomass burning as the dominant source (Seco et al., 2013; Tan et
396 al., 2021). However, the peaks coincided with increases in various other VOCs and pollutants;
397 therefore, the origin of this factor cannot be completely attributed to biomass-burning events.

398

399 At BCN, *Factor 2* has a high correlation with the *traffic & industry* factor (*Factor 1*) (summer: $r = 0.80$;
400 winter: $r = 0.87$) and BC concentrations (summer: $r = 0.82$; winter: $r = 0.79$), which indicates the
401 anthropogenic origin of *Factor 2*. This is further confirmed by the two peaks in the morning and
402 evening, as observed for the *traffic & industry* factor (*Factor 1*), which is common for anthropogenic
403 contaminants at BCN. Notably, the levels of *Factor 2* were much higher during winter when compared
404 to the summer period due to increased anthropogenic contamination, which can also be observed
405 with the higher concentrations of the weekly averages when compared to weekends. During summer,
406 this difference was also much smaller (**Error! Reference source not found.**).

407 Although biomass burning had a major impact on this factor at BCN, it is not likely its sole source.
408 Benzene is emitted from traffic emissions, which cannot be ignored in an urban area such as BCN
409 (Fernández-Iriarte et al., 2020; Gelencsér et al., 1997; Heeb et al., 2000). Furthermore, the diel cycles
410 of acetaldehyde and acetonitrile showed different patterns when compared to this factor (**Error!
411 Reference source not found.**). Acetaldehyde can have various origins, such as fossil fuel combustion
412 (Nogueira et al., 2015; Sinharoy et al., 2019; Xu et al., 2022), biomass burning (Cerqueira et al., 2013;
413 Holzinger et al., 2005, 1999), hydrocarbon oxidation (Luecken et al., 2012), and even vegetation

414 emissions (Seco et al., 2007). At BCN during winter, acetaldehyde showed the highest mixing ratios
415 during midday, with an additional increase in the evening. However, there were no weekly patterns,
416 with similar concentrations observed during weekly averages and weekend averages. The levels
417 during summer were much lower, showing peaks in the morning and evening. The lower
418 concentration during the day could be caused by the fact that acetaldehyde has an atmospheric
419 lifetime of only a few hours (Possanzini et al., 2002; Seco et al., 2013) due to photochemical
420 destruction (Atkinson and Arey, 2003), which resulted in lower mixing ratios being observed during
421 the day in the summer season. This effect was not observed in the winter due to less sunlight and
422 fewer daylight hours, which reduces the ability of acetaldehyde to be removed from the atmosphere
423 when compared to summer (Filella and Peñuelas, 2006; Viskari et al., 2000).

424 Acetonitrile showed higher levels during the day when compared to the evening during both summer
425 and winter at BCN, with higher levels in the winter (Figure 3). Compared to acetaldehyde, the levels
426 during weekends were much lower, indicating an anthropogenic source. Acetonitrile is generally
427 considered a biomass burning marker with a minor source from fossil fuel combustion, which might
428 be the main source at BCN (Holzinger et al., 2005, 1999). Given the stark difference between weekdays
429 and weekends, this is most likely the case.

430 An investigation of the three biggest tracers of *Factor 2* indicates that while this factor partially
431 included biomass burning at BCN, it also was likely a second traffic source in addition to *Factor 1*.
432 Previous studies assessing sources of biomass burning at BCN stated that after the switch to natural
433 gas, the contribution from residential heating in the particulate matter was relatively low (AIRUSE,
434 2016; Amato et al., 2016; Viana et al., 2013). Notably, its contribution remains significant since a study
435 conducted during winter 2011 at BCN attributed 8% of the PM_{2.5} mass to biomass burning (Reche et
436 al., 2012), while a recent study attributed 15% of the total organic aerosols in PM₁₀ to biomass burning
437 (in 't Veld et al. To be Published), indicating a considerable contribution to atmospheric pollution.

438 At MSY, the peaks associated with anthropogenic traffic emissions were absent due to the distance
439 from their source, with the mixing ratios of this factor being mostly even throughout the day. During
440 summer, the diel cycle of *Factor 2* showed a peak at approximately 18:00. (Figure 6); however, this
441 was due to a biomass-burning event observed on the 13th of June. The correlation with BC further
442 indicates that this peak was caused by a biomass burning event, with a correlation of $r = 0.66$ for the
443 entire dataset of *Factor 2* during summer, while the *traffic & industry factor (Factor 1)* did not correlate
444 with either this factor or the BC levels. In winter, this factor likely incorporated pollutants from
445 residential heating and other sources transported from BCN and surrounding industrial areas. This was
446 also confirmed by a correlation with the *traffic & industry factor* ($r = 0.65$), as well as the appearance
447 of a peak at 07:00 for all tracers of *Factor 2*.

448 The diel cycle of benzene at MSY showed similar patterns to the *traffic & biomass factor (Factor 2)*;
449 however, just like BCN, acetaldehyde and acetonitrile had different diel cycles when compared to
450 *Factor 2*. Both compounds showed higher levels in summer than in winter, increasing during the day
451 and declining in the afternoon. In the case of acetonitrile, this could include long-range transport from
452 forest fires in the Mediterranean area (Seco et al., 2013). Notably, acetonitrile levels were significantly
453 lower when compared to BCN, suggesting that acetonitrile at BCN was of local origin.

454 Acetaldehyde also has a biogenic origin, which was more prominent at MSY (Lewis et al., 2005;
455 Luecken et al., 2012; Possanzini et al., 2002). Furthermore, it can also be formed in the atmosphere
456 by photochemical production from alkenes (including biogenic alkenes such as terpenes) (Luecken et
457 al., 2012; Seco et al., 2007). Higher levels in the summer compared to winter could be a result of the
458 increased physiological activity of vegetation and increased secondary production due to greater
459 photochemical activity, as well as higher local acetaldehyde emissions from forest fires (Seco et al.,
460 2011).

461 A previous study by Yáñez-Serrano et al. (2021) near MSY did not identify a biomass burning source
462 during summer due to its low contributions from biomass burning during this season and the lack of

463 confirmed wildfires during their study. Tan et al. (2021) also identified a biomass burning source in
464 Hong Kong using the same tracers as this study.

465 3.3.3 Factor 3: Isoprene Oxidation Source

466

467 *Factor 3* was traced by isoprene (81%) and MVK+MACR (65%), both of which are oxidation products
468 of isoprene (Ling et al., 2019; Liu et al., 2013). This factor also had lower levels of OVOCs attributed to
469 it, such as MEK (15%), acetone (10%), and monoterpenes (10%). The diel cycles of MSY showed a clear
470 increase during daylight in summer, which was less pronounced at BCN (Figure 8). Both stations
471 confirmed this trend with a significant correlation of *Factor 3* with solar radiation (BCN: $r = 0.48$; MSY:
472 $r = 0.52$) at both stations (Figures S1 and S3).

473 At MSY, the diel cycle indicated a biogenic origin for isoprene (Figure 3), which is known to be emitted
474 by vegetation (Pacífico et al., 2011; Sanadze, 2004). This was further confirmed by the high correlation
475 between this factor and the *monoterpene factor (Factor 4)* ($r = 0.91$) during summer, which was also
476 emitted by biogenic activity (see description for Factor 4). This confirms the findings of Yáñez-Serrano
477 et al. (2021a), who also measured high levels of isoprene and its oxidation products in Montseny
478 Natural Park (a few kilometers from our station) during summertime and found a diel cycle similar to
479 the solar radiation cycle, which was also observed in this study ($r = 0.52$). This could be surprising since
480 the forest at MSY is dominated by holm oak, which typically emits isoprene at a lower rate (only 5%
481 compared to monoterpenes) (Peñuelas et al., 2009). Furthermore, several *Quercus* species, such as
482 *Quercus ilex* (a dominant species in the MSY region), also emit isoprene (Fernández-Martínez et al.,
483 2018). However, a previous report for the same MSY site by Seco et al. (2011) found summertime
484 isoprene mixing ratios amounting to approximately two-thirds of those of monoterpenes and with
485 very similar diel patterns, which is in agreement with the present study and suggests that other
486 vegetation present in those forests and the surrounding areas are isoprene emitters. While the exact
487 origin of isoprene cannot be confirmed yet, it was speculated these come from *Erica arborea*, a shrub

488 located near MSY and known to emit a considerable amount of isoprene (Pagès et al., 2020; Yáñez-
489 Serrano et al., 2021a). Yáñez-Serrano et al. (2021a) also hypothesized that isoprene could originate
490 from local vehicle exhaust emissions (Borbon et al., 2001), but no correlation was found between
491 isoprene and any anthropogenic contaminant (Figure S3). Meanwhile, concentrations during winter
492 were considerably lower due to the smaller biogenic emissions as a result of reduced solar radiation
493 and temperature. Moreover, the diel pattern of the isoprene to MVK+MACR ratio (**Error! Reference**
494 **source not found.**) showed that during summer at MSY, the ratios decrease during daylight hours until
495 18:00, with a maximum at approximately 09:00. This indicates that biogenic isoprene becomes
496 oxidized into secondary MVK and MACR.

497 At BCN, a different situation occurred. *Factor 3* showed an increase during daylight hours and
498 decreased in the afternoon, with an additional small peak in the morning and evening (**Error!**
499 **Reference source not found.**). The morning and evening peaks coincide with those from other
500 anthropogenic emissions, which implies there was also an anthropogenic source for this factor (**Error!**
501 **Reference source not found.**). The anthropogenic origin was further confirmed with the positive
502 correlation with BC ($r = 0.50$), the *traffic & biomass burning factor (Factor 2)* ($r = 0.56$), and the *traffic*
503 *& industries factor (Factor 1)* ($r = 0.44$) (Figure S1) during summer. The correlation was even higher
504 during winter, with a positive correlation with BC ($r = 0.46$), *traffic & biomass burning factor (Factor*
505 *2)* ($r = 0.72$), and the *traffic & industries factor (Factor 1)* ($r = 0.60$). Isoprene is known to be emitted
506 by anthropogenic activities, particularly traffic (Borbon et al., 2002, 2001; Filella and Peñuelas, 2006;
507 Wagner and Kuttler, 2014; Yee et al., 2020). However, other anthropogenic VOCs can also contribute
508 to the signal of m/z 69 since it is a common fragment of cycloalkenes (Gueneron et al., 2015).
509 Nevertheless, the maximum isoprene mixing ratios in summer occurred at 13:00, when temperature
510 and solar radiation were also at their maximum. This observation suggests that a biogenic source
511 remains present at BCN, which is logical given the forests surrounding the area.

512 The isoprene oxidation factor was previously identified by Yáñez-Serrano et al. (2021) in Montseny as
513 a biogenic source along with monoterpenes. While monoterpenes are mostly of biogenic origin,
514 isoprene is not necessarily biogenic. Combining the datasets of the two stations (i.e., BCN and MSY)
515 separated this biogenic source between the *isoprene oxidation factor (Factor 4)* and the *monoterpene*
516 *factor (Factor 5)*. The *isoprene oxidation factor* is mostly biogenic at MSY, with high levels in the
517 summer and a clear correlation with temperature (Figure S3). Meanwhile, at BCN, an additional
518 anthropogenic source of isoprene exists, especially during winter, as shown by the diel cycle and
519 correlation with other anthropogenic contaminants (Figure 9; Figure S1).

520

521 3.3.4 Factor 4: Monoterpenes

522 The monoterpene factor (Factor 4) was mostly traced by monoterpenes (84%), with minor
523 contributions from methanol (10%) and C8 aromatics (10%). The monoterpene factor was emitted
524 from biogenic sources at MSY, as seen in the diel pattern (Figure 10). Just like for the isoprene
525 oxidation factor, levels were extremely high during the day while being low at night and in the
526 evening. Furthermore, a clear distinction between seasons is apparent since this factor exhibited its
527 minimum mixing ratios during winter without any diel cycle, whereas it had a clear diel cycle
528 correlating with sunlight and temperature in summer (Figure S3-4). The high levels measured are
529 due to the MSY station being embedded among holm oak trees, which emit these compounds in
530 relevant quantities (Peñuelas et al., 2009; Seco et al., 2013, 2011). This factor was previously
531 identified by Yáñez-Serrano et al. (2021) at Montseny, where it was combined with the isoprene
532 oxidation factor into a single biogenic source.

533 This was not the case for BCN, where the diel cycle did not resemble one increasing with sunlight. This
534 factor did correlate with the anthropogenic factors (*traffic & biomass burning and traffic & industries*)
535 during both seasons, with a correlation factor between 0.61 and 0.72 during summer and 0.71 and
536 0.78 during winter (Figure S1-2). However, the peaks in the morning and evening did not coincide with

537 the peaks of the anthropogenic contaminants (Figures 2 and 10). Hellén et al. (2012) proposed a
538 traffic-related origin of monoterpenes since the diel cycle of monoterpenes resembles those of
539 anthropogenic aromatics; however, as previously mentioned, the peaks at BCN do not coincide with
540 said anthropogenic aromatics. It is more likely that the diel cycle shape was caused by the rapid
541 reaction of monoterpenes with daytime generated OH[•] radicals and dilution by atmospheric mixing,
542 as was reported previously in the study area (Filella and Peñuelas, 2006; Harrison et al., 2001).
543 Although lower levels were observed during weekends, those values might still indicate an
544 anthropogenic origin. The reaction rates with OH[•], presented in Table 6, indicate that the reactivity of
545 monoterpenes is at least 10- to 100-fold faster when compared to the other VOCs.

546 3.3.5 Factor 5: Long-lifetime VOCs

547 The final factor was a mixture of the remaining VOCs in the study area. These were traced by various
548 OVOCs, including acetone (71%), methanol (69%), acetonitrile (61%), MEK (40%), MVK+MACR (31%),
549 and acetaldehyde (23%).

550 Methanol and acetone had the highest mixing ratios at both stations during both seasons, which can
551 be attributed to their long lifetime in the atmosphere and the fact that they are formed from the
552 oxidation of other VOCs (Seco et al., 2007) (Table 6). This resulted in high background levels, as
553 observed in previous studies (Filella and Peñuelas, 2006; Lewis et al., 2005; Seco et al., 2013; Yáñez-
554 Serrano et al., 2021a). Acetonitrile had much lower mixing ratios but also has a long atmospheric
555 lifetime ranging from months to years (Sanhueza et al., 2004). Therefore, this factor was called the
556 *long-lifetime VOCs factor*.

557 At BCN, the *long-lifetime VOCs factor* showed a similar diel pattern to those of methanol and acetone,
558 which was similar to other anthropogenic contaminants during summertime (Figure 11). This resulted
559 in positive correlations with said factors (*traffic & industries*: 0.49; *traffic & biomass burning*: 0.52), as
560 well as the observation of two peaks—one early in the morning and one in the evening—with much
561 lower mixing ratios during weekends. At BCN, Filella & and Peñuelas (2006, and references therein)

562 found that both methanol and acetone were emitted by road traffic, which resulted in peaks during
563 rush hours. The observed decrease of *factor 5* during the day would be a result of photochemical
564 destruction and dilution effects. The anthropogenic origin of these VOCs is further confirmed by the
565 lower mixing ratios over weekends.

566 During winter at BCN, the highest mixing ratios of *Factor 5* occurred at midday, with a small peak in
567 the evening (Figure 11). During summer, the disparity between the weekday and weekend averages
568 was highly accentuated (Figure 11); however, this difference was only observed in the middle of the
569 day during winter. This suggests that at BCN, this factor has a combination of sources—mostly
570 anthropogenic and from atmospheric oxidation—and that this combination can change with the
571 season. The higher levels observed during winter were most likely the result of less sunlight and the
572 consequently diminished photochemical destruction of methanol and acetone.

573 During summer at MSY, the long-lifetime VOCs factor mixing ratio does not represent the trends of
574 any VOCs. This is most likely a result of the different trends that all the VOCs have at this station, which
575 averages to an overall neutral trend. Methanol showed three peaks (Figure 3), with one in the morning
576 (08:00) possibly coinciding with stomatal opening (Fall and Benson, 1996; Hüve et al., 2007), and two
577 more around midday and in the evening. This suggests both the biogenic and photo-oxidation origins
578 of methanol at the MSY station. On the other hand, acetone only displayed a peak during the day that
579 was highly correlated with temperature ($r = 0.75$) and the biogenic sources (isoprene oxidation: $r =$
580 0.69 ; monoterpenes $r = 0.56$). This would indicate that, like methanol, acetone at MSY can be both
581 formed due to photo-oxidation and biogenic emissions at MSY. The only difference was that acetone
582 did not have a stomatal release in the morning, hence the lack of a morning peak.

583 On the other hand, during the winter season, all three long-lifetime VOCs showed no diel cycle due to
584 lower biogenic emissions and sunlight. This indicates that they are mostly long-range transported
585 VOCs due to their lengthy atmospheric lifetimes.

586 3.4 Ozone formation potential

587 The contribution of VOCs to the formation of O₃ was estimated using the MIR method to calculate and
588 assess the contribution of each VOC to ozone formation. Table 5 presents the MIR of each VOC. All
589 MIRs were obtained from Venecek et al. (2018), except for acetonitrile, which has a currently unknown
590 MIR. However, it is important to consider that the MIR values used in this comparison are based on
591 an urban environment, which limits the accuracy of the estimated OFP for MSY (Venecek et al., 2018),
592 although the MIR has been applied in rural environments to estimate the OFP (Gómez et al., 2020;
593 Kumar et al., 2018; Luo et al., 2020). The results of the OFP are presented in Table 5 and Figure 12.
594 The OFP of the sum of the VOCs was similar during summer at both stations, with 18.51 ppbv at BCN
595 and 19.07 ppbv at MSY. During winter, a clear difference was observed. The winter OFP was 16.50
596 ppbv at BCN, only slightly lower compared to the summer season, while the OFP was considerably
597 lower at MSY during winter at 8.42 ppbv. The differences in OFP between the stations and seasons
598 were due to low concentrations of high-MIR VOCs during winter in the rural area. There are two
599 groups with high MIR values among our measured VOCs. These were primarily the aromatics, which
600 are mostly anthropogenic and thus quite low at MSY, with the second-largest group being biogenic
601 VOCs such as isoprene, its oxidation products, and monoterpenes, whose emissions considerably
602 decrease during winter due to a decrease in sunlight. The OFP does not completely account for the O₃
603 concentrations measured simultaneously. At BCN during winter, the OFP was 36% of the total O₃,
604 while it was 13% at MSY during winter. Meanwhile, during summer, BCN exhibited 28% of the total
605 O₃, while MSY exhibited 22%. This dispersion can be due to the limited VOC set measured, as well as
606 various other sources of O₃.

607 Comparing the hydrocarbon VOCs (HVOCs) and OVOCs showed a similar situation during winter at
608 both stations, where OVOCs drove the OFP slightly higher, with 7.22 ppbv (51%) at BCN and 4.89 ppbv
609 (58%) at MSY, compared to 6.93 ppbv (49%) and 3.54 (42%) for HVOC at BCN and MSY, respectively.
610 The situation was quite different during the summer months for the two stations. At BCN, the OFP
611 was mostly driven by HVOCs, with 11.7 ppbv (63%) compared to 6.81 ppbv (37%) of OVOCs. At MSY,

612 the situation was reversed, where the OFP was primarily driven by OVOCs, with 13.06 ppbv (68%)
613 compared to 6.01 ppbv (32%) of HVOCs. This difference in contribution is a commonly observed
614 phenomenon, where OVOCs contribute significantly more to rural areas when compared to urban
615 areas (Chen et al., 2022; Louie et al., 2013; Luo et al., 2011).

616 At BCN, the OFP primarily originates from the *traffic & industry factor (Factor 1)* since both toluene
617 and C₈ aromatics, which were the main tracers of this factor, together made up 50% of the estimated
618 OFP during summer, with 9.13 ppbv. On the other hand, due to the relatively lower MIR and mixing
619 ratios of benzene, it contributed very little toward OFP. The total OFP from the aromatics (benzene,
620 toluene, and C₈ aromatics) was nearly 1.5 times lower during winter (6.28 ppbv) due to the lower
621 mixing ratios of toluene and C₈ aromatics, accounting only for 39% of the total OFP. At MSY, the
622 contribution of the *traffic & industries factor* was significantly lower, with the relative contribution of
623 toluene and C₈ aromatics to OFP being 35% (2.85 ppbv) during winter, which is nearly 2.2-fold lower
624 when compared to BCN during the same season. During summer, the contribution of these aromatics
625 to OFP was only 7% with 1.37 ppbv, representing one of the smallest contributing sources and 6.6
626 times lower when compared to summer at BCN. Although this is worth keeping in mind, since MSY is
627 a rural area, most toluene and C₈ aromatics will be transported from industrial and urban sources and
628 are thus already photo oxidized to O₃ and SOAs when the air masses reach the station (Seco et al.,
629 2013).

630 The biogenic sources, which are the *monoterpenes factor (Factor 4)* and the *isoprene oxidation*
631 *products (Factor 3)*, had much higher mixing ratios during summer due to increased biogenic emissions
632 at both stations. The combination of isoprene, MVK+MACR, and monoterpenes accounted for 20% of
633 the OFP during summer at BCN. This significantly decreased during winter to a combined 7%, with an
634 OFP of 1.16 ppbv compared to 3.71 ppbv during summer. Notably, most of the OFP during winter
635 originated from isoprene and its oxidation products since the monoterpene's contribution at BCN was
636 low. As previously established, this also includes anthropogenic isoprene. At MSY, these sources were

637 mainly biogenic and contributed much more to the total OFP. During summer, the biogenic sources
638 made up 45% of the total OFP with 8.59 ppbv, which was much lower during winter when they only
639 comprised 11% of the total OFP with 0.91 ppbv. This made biogenic VOCs the main driver of OFP at
640 MSY during summer.

641 The final significant contribution to OFP was from acetaldehyde, which, with its high MIR and mixing
642 ratios, contributed significantly to the OFP at both stations. During winter, it contributed 34% of the
643 total OFP at BCN with 5.60 ppbv, which was much higher than during summer, when it only
644 contributed 10% of the total OFP with 1.89 ppbv. During winter at MSY, acetaldehyde also contributed
645 35% of the total OFP, with 2.94 ppbv, while during summer it only contributed 27% of the total OFP
646 but it contributed more (5.07 ppbv) in absolute terms. Ultimately, acetaldehyde contributed relatively
647 the most during the winter season at both stations, even though the actual OFP was higher during the
648 summer season at MSY. Methanol followed closely with 12 to 14% of the total OFP at all stations.

649 3.5 SOA formation potential

650 Most of the existing literature has estimated the SOAP using the relative SOAP to toluene ratio
651 obtained from Derwent et al. (2010). While useful, the list presented there does exclude some
652 essential VOCs that we measured in our study. Therefore, we obtained the SOAP from Gu et al. (2021),
653 who reported the SOA yield of various VOCs included in this study. It is crucial to state that this
654 method, like the OFP, has the limitation that the SOA yield was calculated based on an urban
655 environment with a high NO_x atmosphere, which influences SOA formation (Camredon et al., 2007).
656 While this situation applies to BCN, it is more inaccurate in the case of MSY. Additionally, it is worth
657 mentioning that the SOA yields used OH^\bullet as the primary oxidant to form SOA (Gu et al., 2021). O_3 is a
658 secondary oxidant that can also form SOA, and with O_3 levels being generally higher in the MSY rural
659 area, its capability to oxidize will increase accordingly. However, the reaction rates of VOCs with O_3
660 are 10^5 to 10^{11} times lower (Atkinson and Arey, 2003; Atkinson and Carter, 1984) than those with OH^\bullet ,

661 thereby diminishing the impact of this difference. The results of the SOAP are presented in Table 7
662 and Figure 12.

663 Overall, aromatics contributed the most due to their high SOA yields at both stations during both
664 seasons (Table 7, Figure 12). Since the mixing ratios of these aromatic compounds were much higher
665 at BCN, a much higher SOAP was observed in this study. Due to the aromatic VOCs' high SOA yields,
666 they contributed the most to the SOAP at both stations during both seasons (Table 7). Since the mixing
667 ratios of these aromatic compounds were much higher at BCN, the SOAP was considerably higher
668 compared to MSY. At MSY, the SOAP was similar between the two seasons but was dominated by
669 anthropogenic aromatic compounds (benzene, toluene, C₈ aromatics) during winter, while the
670 monoterpenes were a considerable source of SOAP in summer.

671 This would indicate that *traffic & industry (Factor 1)* was the main contributor to SOA formation from
672 the measured VOCs at BCN since none of the remaining OVOCs and isoprene oxidation products have
673 a significant SOA yield. Although monoterpenes have a significant SOA yield, the levels at BCN were
674 much lower, resulting in minimal influence. However, the diel cycle of the *monoterpene factor (Factor*
675 *4)* suggests that during summer at BCN, monoterpenes were rapidly reacting with daytime-generated
676 OH[•] radicals, which would result in the formation of both O₃ and SOA; therefore, the actual
677 contribution of monoterpenes would be much higher than the values given in Table 7.

678 At MSY during winter, the *traffic & industry factor (Factor 1)* seemed to be the greatest source of SOA
679 in the area. However, as was the case with the OFP, the actual SOAP might be higher as the
680 transported toluene and C₈ aromatics deplete during transport. During summer, the monoterpenes
681 were the main contributors to the SOAP, while other biogenic compounds, such as isoprene, did not
682 contribute as much due to their low SOA yields.

683 4.0 Conclusions

684 In this study, we quantified 11 VOC mixing ratios during summer and winter at both the BCN urban
685 background site and the MSY rural background site (Table 5). Applying PMF to the combined dataset
686 revealed five factors (Table 6). The first factor, *traffic & industries (Factor 1)*, originated from
687 anthropogenic sources such as traffic and industry in and around Barcelona, mostly consisting of aged
688 air masses at MSY. The second factor, *traffic & biomass burning (Factor 2)*, represented a secondary
689 source combining traffic emissions and biomass burning. The third factor, *isoprene oxidation (Factor*
690 *3)*, was associated with isoprene and its oxidation products, which had both anthropogenic and
691 biogenic origins at BCN and were mostly biogenic at MSY. The fourth factor, *monoterpenes (Factor 4)*,
692 was predominantly biogenic at both stations, with lower levels at BCN due to rapid oxidation by OH^{*}
693 and lower biogenic emissions in the area. The fifth factor (*Factor 5*) represented a mixture of VOCs
694 with longer atmospheric lifetimes and included methanol, acetone, and acetonitrile, which had longer
695 lifetimes when compared to other VOCs.

696 Analyzing the sources of the VOCs in relation to OFP revealed that at BCN, anthropogenic VOCs—
697 particularly toluene and C₈ aromatics—were the primary contributors to O₃ formation, accounting for
698 38 and 49% of the measured OFP during winter and summer, respectively (Table 7). BCN also had
699 significant contributions from *isoprene oxidation products* (6 and 19% during winter and summer,
700 respectively) and acetaldehyde (34 and 10% during winter and summer, respectively). Methanol
701 contributed 14 and 13% during winter and summer, respectively. At MSY, biogenic sources—and
702 specifically *isoprene oxidation products* and *monoterpenes*—played a larger role in the total OFP,
703 contributing 11 and 45% during winter and summer, respectively. The aged anthropogenic VOCs from
704 the *traffic & industries factor* made a considerable contribution of 34% during winter and 7% during
705 summer. Acetaldehyde contributed 35 and 27%, while methanol contributed 12 and 14%, during
706 winter and summer, respectively.

707 At BCN, the results demonstrate that both O₃ and SOA formation were predominantly driven by
708 anthropogenic emissions, particularly those originating from the *traffic & industries factor*. The
709 aromatic compounds had a significant impact on the SOAP due to their higher SOA yields (Table 9).
710 The seasonality showed a significant increase during summer, with concentrations of 0.766 μg m⁻³
711 compared to 0.603 μg m⁻³ during winter, primarily due to higher levels of toluene and C₈ aromatics
712 during summer (Table 9). The OFP also showed seasonality with higher levels during the summer, but
713 the difference in total OFP was small, with 16.5 ppbv during winter and 18.51 ppbv during summer
714 (Table 7). The MSY rural background site exhibited a different pattern, where both OFP and SOAP were
715 driven by aged anthropogenic emissions, predominantly from the *traffic & industries factor*, during
716 winter. In contrast, during summer, biogenic sources—specifically the *isoprene oxidation products* and
717 *monoterpenes* factors—played a more prominent role. In the case of SOAP, the *monoterpenes factor*
718 contributed 53% of the overall SOAP during summer, while aromatic compounds contributed 29%.
719 Comparing the total OFP between the two stations, BCN and MSY revealed similar values during the
720 summer season, with 18.51 ppbv at BCN and 19.07 ppbv at MSY. However, during winter, the OFP was
721 considerably lower at MSY, with 8.42 ppbv compared to 16.50 ppbv at BCN. On the other hand, the
722 SOAP exhibited higher concentrations at BCN when compared to MSY, with 1.163 μg m⁻³ and 0.603
723 μg m⁻³ during winter at BCN and MSY, respectively, and 1.833 μg m⁻³ and 0.766 μg m⁻³ during summer
724 at BCN and MSY, respectively.

725 While this study provides valuable insights into the sources of VOCs and their impacts on OFP and
726 SOAP, some knowledge gaps should be acknowledged. First, both OFP and SOAP estimations are
727 based on high nitrogen oxide (NO_x) environments, which may not accurately represent the low NO_x
728 conditions at the MSY rural background station. Secondly, the biogenic contribution of monoterpenes
729 is likely underestimated at BCN during summer due to their rapid oxidation by OH[•] in an urban
730 environment. Lastly, this study focused on a selection of 11 VOCs, and although they make substantial
731 contributions to atmospheric concentrations, the inclusion of additional VOCs—particularly alkanes

732 (~5–10% contribution to OFP and SOAP) and alkenes (>30% contribution to OFP)—would provide a
733 more comprehensive understanding (Jookjantra et al., 2022; Wang et al., 2023; Wu et al., 2017; Yuan
734 et al., 2009; Zhan et al., 2021).

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